Some statistical methods for random process data from seismology and neurophysiology

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Technical Report No. 84
December 1986

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THE 1983 WALD MEMORIAL LECTURES

SOME STATISTICAL METHODS FOR RANDOM PROCESS DATA

FROM SEISMOLOGY AND NEUROPHYSIOLOGY

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"The purpose of statistics,...., is to describe certain real phenomena."

A. Wald (1952)

We present examples of some statistical techniques for the analysis of random process data and examples of their uses in the substantive fields of seismology and neurophysiology. Problems addressed include; frequency estimation for decaying cosinusoids, measuring association, assessing causal connections, estimation of speed and detection, structural modelling. Techniques employed include; complex demodulation, nonlinear regression, probit analysis, deconvolution, maximum likelihood, singular value decomposition, Fourier analysis.

I. INTRODUCTION

The concern of these lectures is raw data distributed in time and/or space. The basic data are curves and surfaces. If n denotes the sample size and p denotes the dimension, then the concern is with the case of $n \ll p$. In the situations addressed, the phenomena have developed or are developing in time or space. They are complex, so that subject matter plays essential roles in the analyses made and in the interpretations and conclusions drawn. There need to be combinations of both physical and statistical reasoning. Indeed, a principal goal of the lectures was to bring out the key role that subject matter plays in the analysis of random process data. A further intention was to show that the fields of seismology and neurophysiology are rich in problems for statisticians, particularly those with some interest in applied mathematics. The work presented involves a mixture of data analysis and structural modelling. The problems discussed are specific, but the techniques employed broadly applicable. The data concerned is of high quality, so that detailed analyses are possible. The material presented consists of personal (collaborative) work and a few success stories of other particular methods. An attempt is made to present problems from a unified point of view.

The study of random process data provides a major interface of statistics with science and technology. Indeed there has been an explosion in the collection of spatial-temporal measurements, (corresponding in part to much of modern

technology having become digital.) Some particular issues and procedures become emphasized as a result of the interaction of statistics with technology. These include: system identification, systems analysis, inverse problems, Fourier inference, bias versus variability (resolution versus precision), averaging functions, dynamics, micro versus macro study. These strains run through the examples presented.

Some provisos are necessary. There is no claim made that the analyses are definitive. What is presented is an overview, rather than specific details. Further there is little presentation of formalism.

There are two lectures. The first concentrates on some statistical methods in seismology, the second on some corresponding methods in neurophysiology. It is interesting to see the same methods playing central roles in the analysis of data from two quite disparate fields.

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III. SEISMOLOGY

"Jeffreys ... attention to scientific method and statistical detail has been one of the main forces through which Seismology has attained its present level of precision."

Bullen and Bolt (1985)

1. The Field and Its Goals. The term seismology refers to the scientific investigation of earthquakes and related phenomena. It has been defined as the "science based on data called seismograms, which are records of mechanical vibrations of the Earth", Aki and Richards (1980). This second definition allows the admission that seismologists also study vibrations caused by the sea, by volcanoes or by man. One further definition that has been given is: the science of strain-wave

propagation in the Earth.

Whatever the definition, the broad goals of seismology are to learn the Earth's and planet's interior composition and to predict the time, size, location and strength of ground motion in future earthquakes. Workers in the field seek to provide valid explanations of earthquake related phenomena and to understand those phenomena so that life may be made safer.

Specific problems addressed include: the detection, location and quantification of earthquakes, the distinguishing of earthquakes from nuclear explosions, and the determination of wave velocity in the Earth's interior as a function of depth.

The accumulation of knowledge in seismology has displayed a steady back-and-forth between new insight concerning the waves and new insight concerning the media through which the waves propagate. Among major "discoveries" one can list: the inner core, the liquid central core, the Mohorovic discontinuity, the movement of tectonic plates causing earthquakes themselves, and the locating of numerous gas and oil fields.

The field is largely observational, with the basic instruments the seismogram and clock. There are important experiments too, where tailored impulses are input to the Earth and the resulting vibrations studied. The field experienced the "digital revolution" in the fifties and now poses problems exceeding the capabilities of even today's computers.

Statistical methods have played an important role in seismology for many years - in large part due to the efforts of Harold Jeffreys (see Jeffreys (1977), for example.) Vere-Jones and Smith (1981) provide a review of many contemporary instances. Statistics enters for a variety of reasons. The data sets are massive. There is substantial inherent variability and measurement error. Models need to be refined, fit and revised. Inverse problems need to be addressed. Experiments need to be designed. Sometimes the researcher must fall back on simulations. The basic quantity of concern is often a (risk) probability. In particular, it may be pointed out, that in the construction of the Jeffreys and Bullen (1940) travel time tables, one has an early, perhaps greatest success, of the use of robust/resistant methods. (B. A. Bolt's (1976) presidential address "Abnormal seismology" is well worth reading in this connection.)

Seismologists deal with data of a variety of types. The important forms are digital wave forms from spatial arrays of seismometers of various dimensions (where the instruments have been arranged in such a fashion that an earthquake signal may be seen as a moving, changing shape) and catalogs (containing lists of event's times, locations, sizes and other characteristics) for geographic regions of interest.

Seismology is not without its controversies. There are fundamental ones, such as whether or not plate tectonics is a validated theory. There are practical ones, such as does the size of the motion of an earthquake increase steadily as one approaches the fault or does it level off? As is so often the case, the existing data prove inadequate to resolve these disputes.

A general reference that provides much of the pertinent seismological background is Bullen and Bolt (1985). We turn to a presentation of some specific problems and techniques.

2. Free Oscillations of the Earth. This subject is one of the principal developments in seismology over the last twenty five years. Whenever there is a great earthquake, the Earth vibrates for days afterwards. The seismogram then consists, approximately, of a sum of an infinite number of exponentially decaying

cosinusoids plus noise, see expression (2) below. The frequencies of the cosinusoids and the corresponding rates of decay, relate to the Earth's composition. Measured values may be used to make inferences re that composition. The techniques of complex demodulation, nonlinear regression and regularization may be employed in this connection. Some details on these techniques will follow.

As is the case with many natural systems, the vibratory motion of the Earth may be described by a system of equations of the form

$$\frac{d\mathbf{Y}(t)}{dt} = \mathbf{A}\mathbf{Y}(t) + \mathbf{X}(t) \tag{1}$$

with $\mathbf{X}(.)$ a (vector-valued) input. In the case that the input is $\mathbf{b}\delta(t)$, with $\delta(.)$ the Dirac delta function (corresponding to the earthquake shock) and initial conditions $\mathbf{Y}(0-) = 0$, the general solution of (1) may be written

$$\mathbf{Y}(t) = \exp\{\mathbf{A}t\}\mathbf{b}$$

$$= \sum_j \varsigma_j \exp\{\mu_j t\}\mathbf{u}_j$$

$t > 0$, where μ_j, $\mathbf{u}_j$ are the (assumed distinct) latents of the matrix $\mathbf{A}$. The spectrum occurring is discrete because of the finiteness of the Earth as a body. Focussing on one of the coordinates of $\mathbf{Y}(t)$, and assuming the presence of noise one has

$$Y(t) = \sum_k \alpha_k \exp\{-\beta_k t\}\cos(\gamma_k t + \delta_k) + \epsilon(t) \tag{2}$$

with $-\beta_k$ and γ_k the real and imaginary parts of the μ_j and $\epsilon(.)$ the noise. This model may be checked by complex demodulation of the series $Y(t)$ in the neighborhood of frequencies β_k, as estimated from the periodogram. Provided the bandwidth of the demodulation is not too great a single cosinusoid should be included, the log amplitude should fall off linearly and the phase angle should be approximately constant. Details are given below, specifically at (5).

Figure 1 is a plot of the seismogram recorded at Trieste of the 1960 Chilean great earthquake after removing the tides. Figure 2, the logarithm of the periodogram of this data suggests the presence of many periodic components. The periodogram of a stretch of time series values $Y(t)$, $t = 0,...,T-1$ is defined as follows; set

$$d_Y^T(\lambda) = \sum_{t=0}^{T-1} Y(t)\exp\{-i\lambda t\} \tag{3}$$

$-\infty < \lambda < \infty$, then the periodogram at frequency λ is defined as

$$I_{YY}^T(\lambda) = (2\pi T)^{-1}|d_Y^T(\lambda)|^2 \tag{4}$$

For data from the model (2), $I_{YY}^T(\lambda)$ may be expected to show peaks for λ near the γ_k.

The basic ideas of complex demodulation are: frequency isolation by narrow band filtering to focus on a single term in expression (3), followed by frequency translation to slow the oscillations down. The specific steps are: i) $Y(t) \rightarrow Y(t)\exp\{i\lambda t\}$, (modulation), followed by ii) local smoothing in t of $Y(t)\exp\{i\lambda t\}$ to obtain $Y(t,\lambda)$, the complex demodulate at frequency λ. In the case that $Y(t) = \alpha\exp\{-\beta t\}\cos(\gamma t + \delta)$, one has

$$Y(t,\lambda) \approx \frac{1}{2}\alpha e^{i\delta}e^{-i\beta t}e^{i(\lambda-\gamma)t} \quad \text{for } \lambda \text{ near } \gamma \tag{5}$$

$$\approx 0 \quad \text{otherwise}$$

Hence $\log |Y(t,\lambda)| \approx \log \frac{\rho}{2} - \beta t$ and $arg\{Y(t,\lambda)\} \approx \delta+(\lambda-\gamma)t$. Plots of these quantities, versus t, provide checks on model adequacy and provide preliminary estimates of parameters. Figures 3 and 4 presents such plots for the Chilean data for two frequencies, 3.89 and 5.67 cycles/hour. The rate of decay, β, generally varies with frequency in the present seismological situation. Results for the Chilean data for a variety of frequencies may be found in Bolt and Brillinger (1979).

The parameters could be estimated from the complex demodulate pictures. It is generally more effective to proceed via nonlinear regression. This has the further advantage of providing estimated standard errors. Suppose one has a model

$$Y(t) = S(t;\theta)+\epsilon(t)$$

with $S(.)$ known up to the finite dimensional parameter θ and $\epsilon(.)$ a noise series. In the present case $S(t) = \alpha\exp\{-\beta t\}cos(\gamma t +\delta)$ and $\theta = \{\alpha,\beta,\gamma,\delta\}$. For the next step, it is convenient to take $\lambda_j = 2\pi j/T$ and to write $Y_j = d_Y^T(\lambda_j)$, $E_j = d_\epsilon^T(\lambda_j)$ and $S_j(\theta) = d_S^T(\lambda_j)$. One will estimate θ by minimizing

$$\sum_{j \text{ in } J} | Y_j - S_j(\theta)|^2 \qquad (6)$$

for J a range of subscripts with λ_j near γ. The logic of this is the following: there are a variety of central limit theorems for empirical Fourier transforms (see for example Brillinger (1983).) Suppose that the noise series $\epsilon(.)$ is stationary and mixing with power spectrum $f_{\epsilon\epsilon}(\lambda)$. Then for large T, E_j is approximately complex normal with mean 0 and variance $2\pi Tf_{\epsilon\epsilon}(\lambda_j)$. Further the variates E_j, E_k are approximately independent. It follows that the determination of an estimate of θ to minimize expression (6) is approximately the maximum likelihood procedure. The statistical properties of such estimates were indicated in Bolt and Brillinger (1979) and developed in detail in Hasan (1982). For example one finds the asymptotic variance of $\hat{\gamma}$ to be proportional to

$$\frac{4\pi f_{\epsilon\epsilon}(\gamma)}{T^3\alpha^2}$$

having considered a limiting process with $\beta=\phi/T$ as $T\to\infty$.

Earlier in the paper it was noted that progress in seismology shows a to-and-fro between new knowledge of waves and new knowledge of the structure of the Earth. This occurs in the case of free oscillations. Suppose one has an initial model for the Earth in terms of some physical parameters, eg. expressions for density, shear wave velocity and compression wave velocity as functions of depth, say $\rho(r)$, $c_S(r)$, $c_P(r)$ respectively, r denoting depth. Figure 5, based on the data in Tables 3 and 4 of Bolt (1982), shows what is meant by an Earth model. Given such a model, one can compute the implied frequencies of free oscillation, γ_k. How to do this is described in Chapter 6 of Lapwood and Usami (1981) for example. The relationship involved is nonlinear, but perturbations may be expressed linearly via kernels. Specifically, suppose one perturbs the parameters by amounts $\Delta\rho$, Δc_S, Δc_P respectively, then the perturbation of the frequency of the k-th free oscillation is given by

$$\Delta\gamma_k \approx \int_0^R A_k(r)\Delta\rho(r)dr + \int_0^R B_k(r)\Delta c_S(r)dr + \int_0^R C_k(r)\Delta c_P(r)dr \qquad (7)$$

for kernels A_k, B_k, C_k. Expression (7) is said to lay out the "direct problem" : given $\Delta\rho$, Δc_S, Δc_P find $\Delta\gamma_k$. Now suppose a great earthquake occurs. Then new estimates of the frequencies, γ_k, are available. One has the "inverse problem" :

given the observed $\Delta \gamma_k$, find the $\Delta \rho, \Delta c_S, \Delta c_P$. It seems worth setting out the type of problem involved here in a specific notation. Let Υ and Θ denote normed spaces. Let $\mathbf{X}$ denote a map from Θ to Υ , $\mathbf{Y} = \mathbf{X}\theta$. The values $\mathbf{Y}$ and $\mathbf{X}$ are given, a value for θ is desired. Let α denote a scalar. Some methods for selecting a θ currently being studied include: a) regularization, chose θ to minimize $\|\mathbf{Y} - \mathbf{X}\theta\|^2 + \alpha\|\theta\|^2$, b) sieve, choose θ subject to $\|\theta\| \leq \alpha$ to minimize $\|\mathbf{Y}-\mathbf{X}\theta\|$, c) residual, choose θ subject to $\|\mathbf{y}-\mathbf{X}\theta\| \leq \alpha$ to minimize $\|\theta\|$. A characteristic of the solutions obtained is that one has to be content with the estimation of some form of average of the unknown θ. Chapter 12 of Aki and Richards (1980) contains a discussion of inverse problems in geophysics. A characteristic that distinguishes the present earth model problem, from the usual inverse problems, is that there are discontinuiuties present in the model - corresponding to the Earth's layers.

Several other references to the study of free oscillations may be noted. Hansen (1982) extends the procedure of Bolt and Brillinger (1980) to handle the case of several eigenfrequencies present in the nonlinear regression fit. Dahlen (1982) sets down the asymptotic results for the case of tapered data, that is when convergence factors have been introduced into the Fourier transform computations. Zadro and Caputo (1968) look for nonlinearities via bispectral analysis.

3. Estimation of Fault Plane Parameters The see-saw between the study of the Earth's structure and the study of earthquake sources was pointed out earlier. In this section it will be indicated how a (nonlinear) probit analysis may be employed to estimate basic characteristics of the source of an earthquake.

An important quantity read off the seismic trace of an earthquake at a particular observatory is the sign of the increment at the arrival of the first energy from the event. This sign corresponds to whether the initial motion is a compression or a dilation. In many cases, following the observation of an earthquake at a number of stations, if the observed signs of first motion are plotted on a map centered at the epicenter of the event a (radiation) pattern results. Figure 6, taken from Brillinger, Udias and Bolt (1980), provides such a plot for one of the aftershocks (event 4) of the Good Friday 1964 Alaskan event. Following Byerly (1926), such plots have been employed to learn about the source. Before describing what may be learned, some details of the earthquake process will be set down. The usual assumption (the elastic rebound theory) is that earthquakes are due to faulting. A crack initiates at a point and (in the case of pure slip) spreads out to form a fault plane. As the crack passes a given point, slip takes place (on the fault plane) resulting in a stress drop and the radiation of seismic waves. The radiated (P-) waves may be shown to have a quadrantal pattern with one of the axes parallel and the other perpendicular to the fault plane of the event. It follows, and this is what Byerly (1926) contributed, that the data may be used to estimate the fault plane orientation. Having an estimate of the fault plane and the direction of motion on that plane is important to geology and geophysics. Researchers seek to tie together surface and subsurface features, to consider regional stress directions and to use the results to confirm and extend the theory of plate tectonics. The results can be crucial to seismic risk computations.

Byerly proceeded graphically and this has continued to generally be the working approach. However the results so obtained are subjective, have no attached measure of uncertainty and may not be easily combined with estimates derived from other events at the same site, however.

The problem may be approached in formal statistical fashion as follows: the data available consists of; hypocenter of earthquake, locations of observatories, directions of observed first motions (compressions or dilations) at the observatories and a store of knowledge concerning the Earth's structure (velocity models

as given in Jeffreys and Bullen (1940) for example). It may further be argued that seismographic noise is approximately Gaussian (see Haubrich (1965).) Let a fault plane be described by three angles $(\theta_T, \phi_T, \theta_P)$. Let $A_{ij}(\theta_T, \phi_T, \theta_P)$ denote the theoretical expression for the wave amplitude on the focal sphere for event i at station j. This expression may be found in Brillinger, Udias and Bolt (1980). (The focal sphere is a "little" sphere of unit radius around the hypocenter. In carrying out the amplitude computation one has to trace the ray from the hypocenter to the observatory through the focal sphere.) Let Y_{ij} denote the realized amplitude of the seismogram at the onset of the event. Then one can write $Y_{ij} = \alpha_{ij} A_{ij} + \epsilon_{ij}$, with α_{ij} a scale factor and ϵ_{ij} normal mean 0 and variance σ^2 variate. Let $y_{ij} = 1$ if $Y_{ij} > 0$ and $= 0$ otherwise. It follows that

$$Prob \{y_{ij} = 1\} = Prob \{Y_{ij} > 0\} = \Phi(\rho_{ij} A_{ij})$$

where $\rho = \alpha/\sigma$. The model may be further expanded by including a term, γ_{ij}, to allow for reader and recorder errors now writing

$$Prob \{y_{ij} = 1\} = \gamma_{ij} + (1-2\gamma_{ij})\Phi(\rho_{ij} A_{ij})$$

Precise data correspond to γ and σ small (hence ρ large) and imprecise to γ near .5 or ρ near 0.

The model is seen to take the form of a nonlinear probit, (with a term, γ, corresponding to "natural mortality".) An example of a corresponding likelihood is provided by

$$\prod_{ij} \Phi(\rho_i A_{ij})^{y_{ij}} \left(1 - \Phi(\rho_{ij} A_{ij})\right)^{1-y_{ij}} \tag{8}$$

assuming ρ to depend on event alone and $\gamma = 0$. One can now proceed to estimate the unknown parameters, θ_T, ϕ_T, θ_P, ρ, by maximum likelihood.

Figure 6 includes the fitted planes for the case of event 4 of the Alaska sequence. These particular estimates were computed restricting the likelihood (8) to the observations of event $i = 4$. It is crucial to assess the fit of the model. In Brillinger, Udias and Bolt (1980) this was done by comparing the theoretical and estimated probability functions. Figure 7 is based on a pooled analysis of some 16 of the Alaskan events that seemed to go together. The figure provides the empirical probability that the observed first motion agrees with the theoretical as a function of amplitude. The fitted amplitudes have been grouped into cells of width .1 in the work. What is plotted is

$$\# \{(i,j) \mid sgn \ Y_{ij} = sgn \ \hat{A}_{ij}, A-h < \hat{A} < A+h\}/\# \{(i,j) \mid A-h < \hat{A} < A+h\} \tag{9}$$

for $h = .05$. The fit seems adequate.

The results of further computations of this type may be found in the just cited reference and in Buforn (1982). The maximization program VA09A of Harwell was found effective in determining the maximum likelihood values. The estimates were however non-unique and poorly determined in the cases of some small data sets.

An important byproduct of such analyses is to form clusters of like fault plane solutions for events in the same region, in order to get at motions occurring on the same fault plane. The maximum likelihood standard errors are useful in this connection. The practical implication of the work just reported is that large collections of events may be handled routinely and that geophysical conjectures may be checked formally. The final fault-plane solution may be plotted in traditional fashion allowing examination of the data for difficulties. What remains is for more realistic seismic source models than the one treated in the papers listed, to be fit statistically. An elementary reference to the subject matter of concern

here is Boore (1977).

4. Quantification of Earthquakes. One of the important and difficult questions of seismology is how to measure the "size" of an earthquake. Size is an essential feature that a seismologist makes use of in attempts to deal with earthquake hazards and to understand the basic phenomena of concern. Specifically, the seismologist is not only interested in estimating the direction of movement at the source, he is further interested in the overall deformation that took place and the amount of energy that was released. Among the physical quantities of interest for a given earthquake are the seismic moment (a measure of the seismic energy released from the entire fault) and the stress drop (difference between the initial and final stress.)

For a variety of seimic source models, seismologists have related the seismic moment and stress drop to characteristics of the amplitude spectrum, $|S(\lambda)|$, the modulus of the Fourier transform of the signal. Suppose that the seismogram is written

$$Y(t) = s(t;\theta) + \epsilon(t) \tag{10}$$

where $s(.)$ is the signal, θ is an unknown parameter and $\epsilon(.)$ is a noise disturbance. If $S(\lambda;\theta)$ denotes the Fourier transform of $s(t;\theta)$, then what is given, from the source model, is the functional form of $|S(\lambda,\theta)|$. Common forms (for displacement measurements) include

$$|S(\lambda;\theta)| = \alpha/\sqrt{1 + (\lambda/\lambda_0)^\beta} \quad \text{and} \quad \alpha/\left(+ (\lambda/\lambda_0)^2\right)$$

with $\theta = \{\alpha,\beta,\lambda_0\}$. Estimates of the seismic moment and stress drop may be determined once estimates of α and λ_0 are available. The practice has been to estimate the unknowns graphically from a plot of the modulus of the amplitude of the empirical Fourier transform, $|d_Y^T(\lambda)|$. The following formal procedure was suggested in Brillinger and Ihaka (1982).

The asymptotic distribution of $|d_Y^T(\lambda)|$ may be evaluated in the case of stationary ϵ using a central limit theorem of the type mentioned in Section 2. The asymptotic distribution is found to depend on $|S(\lambda;\theta)|$ and $f_{\epsilon\epsilon}(\lambda)$ alone. Hence one needs an expression only for the modulus of S. Next, with the model (10), and small noise,

$$|d_Y^T(\lambda)| = |S(\lambda;\theta)| + (d_\epsilon^T(\lambda) + d_\epsilon^T(-\lambda))/2 + \ldots$$

showing variation around $|S|$ not depending on $|S|$. However when deviations of $|d_Y^T|$ from a heavily smoothed version of itself are plotted versus the smoothed values, dependence of the error on $|S|$ is apparent. An example is provided in Figure 8. This is the result of computations for an earthquake of magnitude 6.7 that occurred in Taiwan on 29 January 1981. The data was recorded by one of the instruments of the Smart 1 array, see Bolt *et al* (1982). The top graph of the figure provides the transverse S-wave portion of the recorded accelerations. The lower graph provides the deviations plot just referred to. This plot suggests that the noise is in part "signal generated". There are various physical phenomena that can lead to signal generated noise. These include multipath transmission, reflection and scattering. The following is an example of a model that includes signal generated noise.

$$Y(t) = s(t) + \sum_k \left(\gamma_k s(t-\tau_k) + \delta_k s^H(t-\tau_k)\right) + \epsilon(t)$$

with the τ_k time delays, with s^H the Hilbert transform of s and with γ_k, δ_k reflecting the vagaries of the transmission process. The inclusion of the Hilbert

transform allows the presence of phase shifts. With the γ_k, δ_k, τ_k random, and after evaluating the large sample variance, one is led to approximate the distribution of $Y_j = d_j^T(\lambda_j)$ by a complex normal with mean $S(\lambda_j;\theta)$ and variance $\Gamma_j = 2\pi T\{\rho^2 | S(\lambda_j;\theta)|^2 + \sigma^2\}$, where now ϵ has been assumed to be white noise (of variance σ^2), and also it is assumed that $E\gamma_k$, $E\delta_k = 0$, and that the process τ_k is Poisson. The ratio ρ^2/σ^2 measures the relative importance of signal-generated noise. One can proceed to estimate θ by deriving the marginal likelihood based on the $|Y_j|$. This likelihood may be evaluated and found to be

$$\prod_j \left[\exp\left(-\frac{|Y_j|^2 + |S_j|^2}{\Gamma_j} \right) I_0\left(\frac{2|Y_j||S_j|}{\Gamma_j} \right) \frac{1}{\Gamma_j} \right]$$

where I_0 denotes a modified Bessel function. Figure 9 shows a fit of the model $|S(\lambda)| = \alpha|\lambda|/(1+(\lambda/\lambda_0)^4)$ to the data of Figure 8.

Once estimates of α, λ_0 are at hand, these may be converted to estimates of the seismic moment and stress drop via theoretical relationships developed by geophysicists. Details may be found in Ihaka (1985). We remark that this model fit corresponds to a time domain pulse $s(t) = \alpha\lambda_0^2 p(\lambda_0 t)$ where

$$p(t) = [sin\frac{t}{\sqrt{2}} + t\ cos(\frac{t}{\sqrt{2}}-\frac{5\pi}{4})]e^{-t/\sqrt{2}}$$

for $t > 0$ and $p(t) = 0$ otherwise. The form of this pulse indicates how λ_0 corresponds (inversely) to the duration of the event and how α corresponds to size.

5. Array Data.
Today it would be a very strange thing for an earthquake to be recorded on just one seismometer. Indeed from the earliest days, readings of the same event at geographically scattered observatories have been made use of. Since the sixties seismometers have been deliberately arranged in geometric designs over distances of miles to hundreds of miles in order to allow extraction of traditional information and sometimes extraction of new information.

An important use has been the estimation of the direction from which a seismic signal is arriving and the velocity with which it is moving. One manner in which this is done is by the computation of estimates of frequency-wavenumber spectra. The procedure may be described as follows. Suppose one has array data; $Y(x_j,y_j,t)$, $j=0,...,J$, $t=0,...,T-1$. Here (x_j,y_j) denotes the coordinates of the location of the j-th sensor. The frequency-wavenumber periodogram of this data is given by

$$I_Y^T(\mu,\nu,\lambda) = |\sum_j\sum_t Y(x_j,y_j,t)exp\{-i(\mu x_j+\nu y_j+\lambda t)\}|^2 \tag{11}$$

$-\infty < \mu,\nu,\lambda < \infty$. A motivation for this definition is the following. Suppose one has a plane wave $Y(x,y,t) = \rho\ cos(\alpha x+\beta y+\gamma t+\delta)$ of temporal frequency γ and wave number $\kappa = (\alpha,\beta)$. Then the periodogram will have a peak near (α,β,γ). (Incidentally, this wave is moving with apparent velocity $\gamma/\sqrt{\alpha^2+\beta^2}$ from azimuth given by $\tan\phi = \beta/\alpha$.) An example of array data is given by Figure 10. What is plotted are the locations ("*") of 9 of the seismometers of the Smart 1 array located in Taiwan. Also plotted are the portions of the traces used in the computations. These traces correspond to the vertical P-wave part, of the 29 January 1981 earthquake. (The initial near flat part is the noise, saved in a buffer, just before the onset.) The estimated epicenter of this earthquake was 30 km southeast of the array. Figure 11 gives central portion of the frequency-wavenumber periodogram, for this data, as computed via formula (11), at frequency λ corresponding to 1.944 cycles/second. There is seen to be a large peak in

the southeast quadrant, at an azimuth that turns out to correspond to that of the epicenter of the event. The radial distance corresponds to the velocity of P-waves.

Seismologists working with this type of data have often preferred to employ, what they call, the "high-resolution" or "Capon" statistic, see Capon (1969), instead of the periodogram (11). The high-resolution statistic typically shows more dramatic peaks than the periodogram. Before defining it, we introduce some notation. Let $\mathbf{Y}(t)$ denote the j-vector, $|Y(x,,y,,t|$. Set

$$\mathbf{Y}_k = T^{-1} \sum_{t=0}^{T-1} \mathbf{Y}(t) \exp\{-i\frac{2\pi kt}{T}\}$$

for $k = 0,1,2,\dots$. Further let $\mathbf{B} = |exp\{-i(\mu x, +\nu y,)\}|$. If $\lambda = 2\pi l/T$, l an integer, then the periodogram (11) is proportional to $|\bar{\mathbf{B}}'\mathbf{Y}_l|^2$. Next define

$$\mathbf{M} = \sum \mathbf{Y}_k \bar{\mathbf{Y}}_k'$$

with the sum over k with $2\pi k/T$ near λ. Now the high-resolution statistic at frequency λ may be written $1/\bar{\mathbf{B}}'\mathbf{M}^{-1}\mathbf{B}$. If $Y(x,y,t) = \rho cos(\alpha x + \beta y + \gamma t + \delta) + noise$, this statistic may be expected to show a peak for (μ,ν) near (α,β) and λ near γ. This statistic has been introduced, in part, in order to be able to present the next example.

Figure 12 is reproduced from Scheimer and Landers (1974). It shows the high-resolution statistic computed for two portions of data recorded by the Large Apperature Seismic Array in Montana following a strip-mining blast. These computations confirmed the validity of the high-resolution approach. The statistic for one portion shows a single large peak in the direction of the blast. The statistic for the following portion shows energy arriving from various directions. This analysis provided empirical proof of the existence of scattering of seismic waves. That this phenomenon existed had been theorized for years. A frequency-wavenumber data analysis has provided the confirmation.

Spectral analyses are (too) often thought of as being appropriate only for stationary data. As the preceding example shows, the technique may be dramatically useful in nonstationary cases as well. As a second example we mention the results of Bolt *et al* (1982). If, in fact, an earthquake is caused by faulting then the direction of the source of seismic energy will be changing as the fault is ripping, that is as the fault tip is advancing. In the paper cited, Bolt *et al* present high-resolution spectra for succeeding time stretches of the 29 January 1981 Smart 1 event. There is an apparent shift in direction with time. Their work may have been the first experimental measurement of a seismic dislocation moving along a rupturing fault.

In each of the preceding two examples, frequency-wavenumber analysis has allowed researchers to confirm the presence of suspected scientific phenomena.

6. Exploration Seismology (Reflection Seismology). The problem of learning the Earth's crustal structure can be approached as one of system identification. The approach to be described takes advantage of the fact that the Earth is made up of layers. Signals, such as powerful impacts or explosions, can be deliberately input to the Earth and the consequent vibrations recorded by an array of seismometers or geophones. Such experiments may be carried out in a search for gas and oil, or in a scientific study of the general geological makeup of a region of interest. The results of these experiments may be viewed as one of the grand success stories for statistical techniques generally, and of least squares particularly. An unusual aspect of the inferences made, is that in many cases one gets to examine their validity, by the later drilling of a well.

In its simplest form, the energy of an initiated seismic disturbance propagates through the Earth with a spreading wavefront. When it meets an interface between geological strata, part of the energy may be reflected back and part continue forward. The sensors record the returning reflected energy echoes. Knowledge of subsurface velocities allows estimation of the depths and angles of inclination of the various reflectors. Knowledge of the locations of reflectors allows estimation of velocities. In practice the initiating impacts will be repeated a number of times at the same location, and at points of a grid.

If the input signal is taken to be $X(t)$ and if $Y(t)$ denotes the corresponding output, then the two may be modelled as related, assuming linearity and time invariance, by

$$Y(t) = \int a(t-s)X(s)ds \qquad (12)$$

The function $a(.)$ is called the impulse response, since if the Dirac delta function $\delta(t)$ is taken as input, then the resulting output is $Y(t) = a(t)$. The function $a(.)$ evidences the reflectors and velocities in the earth beneath the source and receiver. The model and its interpretration may be motivated as follows. Suppose a pulse is applied at time τ. Suppose in consequence a wave is generated, travels at velocity v_1 to a reflector at distance d_1 and a proportion α_1 is reflected back. With $X(t) = \delta(t-\tau)$, then $Y(t) = \alpha_1\delta(t-\tau-2d_1/v_1)$. (This is actually the naive model for radar or sonar.) Suppose further that the transmitted portion continues downward at velocity v_2 to a reflector at distance d_2 and a proportion of its energy is reflected back, some of which is transmitted by the first reflector to reach the receiver. Now the response has the form $Y(t) = \alpha_1\delta(t-\tau-2d_1/v_1) + \alpha_2\delta(t-\tau-2d_1/v_1-2d_2/v_2)$. This last is seen to correspond to the system of expression (12) with impulse response $a(t) = \alpha_1\delta(t-2d_1/v_1)+\alpha_2\delta(t-2d_1/v_1-2d_2/v_2)$. One can clearly extend this model to situations with many layers, many velocities and many corresponding transmission and relection coefficients. Peaks in the function $a(t)$ may be seen as corresponding to reflectors. (It must be noted that unfortunately such an elementary interpretation is likely to be complicated in practice by interfering phenomena such as ghost reflections unfortunately.) Two references discussing the basics of exploration seismology are Wood and Treitel (1975) and Robinson (1983).

The problem has now been formulated as one of system identification; given stretches of corresponding input, X, and output, Y, determine an estimate of the impulse response $a(.)$. In the case that a pulse close to a Dirac δ may be generated and that the function $a(.)$ drops off to 0 reasonably quickly, a convenient procedure results from taking

$$X(t) = \sum_{m=1}^{M} \delta(t-m\Delta)$$

as input and the "average evoked response"

$$\hat{a}(s) = \frac{1}{M} \sum_{m=1}^{M} Y(s+m\Delta)$$

as an estimate of $a(s)$. This input corresponds to applying pulses periodically.

In practice, the Dirac δ function is an unachievable input and one seeks for feasible probe functions and estimation procedures. In this connection, suppose one sets $m_{YX}(t) = Y*X(t)$ for some convolution operation "$*$". Then from equation (12) one has

$$m_{YX}(t) = \int a(t-s)m_{XX}(s)ds$$

and one has a deconvolution (or inverse) problem to solve. Suppose one decides to seek an $X(.)$ such that

$$\int a\,(t-s)m_{XX}(s)ds \;\approx\; a\,(t)$$

to allow elementary processing. In terms of Fourier transforms, the left hand side here may be expressed as $\int \exp\{i\lambda t\}A(\lambda)M_{XX}(\lambda)d\lambda$, with M_{XX} a Fourier transform of m_{XX}. Then what is wanted is an X such that $M_{XX}(\lambda)\approx 1$ on the support of $A(.)$. If $A(\lambda)$ is known to be near 0 for $0 < \lambda < \lambda_0$ and for $\lambda > \lambda_1$, then a possible function is the "chirp" signal

$$X(t) \;=\; cos\left(|\lambda_0+(\lambda_1-\lambda_0)\frac{t}{\tau}|t\right)$$

for $0 \leq t \leq \tau$. In the seismic case, the values of λ_0, λ_1 have been determined in various experiments. The chirp probe originated in radar work in the second world war, (see Cook and Berenfield (1976).) It may be seen to attach near equal power to the frequencies between λ_0 and λ_1. In the seismic case special devices have been developed to input the chirp signal to the earth. The signal is input repeatedly and the results averaged. The response is then convolved with the chirp function, that is m_{YX} is formed, to estimate $a(.)$. Structure can appear dramatically during the cross-correlation processing described here.

In practice subtle further processing is employed to handle wavefront curvature, ghost reflections, and other natural phenomena that may be present.

7. Other Problems. There are other problems arising in seismology to which statistical methodology can be applied fruitfully. These include: analysis of the coda, analysis of scattering, risk analysis, nonlinear phenomena, point process studies, polarization, cepstral analysis, discrimination of earthquakes from explosions, seismicity study, travel time table construction, attenuation laws, earthquake location, azimuthal dependence of characteristics. Vere-Jones and Smith (1981) discuss several of these problems. In some cases work has begun.

8. Discussion. Seismologists have long been serious users of statistical methods. One finds Harold Jeffreys making the following statement in the entry 'Seismology, statistical methods', in the International Dictionary of Geophysics: *"The uncertainty is as important a part of the result as the estimate itself. ... An estimate without a standard error is practically meaningless."* Hudson (1981) remarks: *"The success of the Jeffreys-Bullen travel time tables was due in large part to Jeffreys' consistent use of sound statistical methods."* When I asked my colleague B. A. Bolt what he saw as the role of statistics in seismology, he replied: *"Seismology is largely an inferential science. ... The role of statistics in seismology is to provide a rigorous procedure for turning observations on seismic waves, etc., into probabilistic statements about properties of the (real) Earth."*

One may note that work in seismology is characterized by: massive data sets, inherent variability and measurement error, defining/fitting/refining models, design of experiments, simulation, probabilistic description, needs for robust/resistant procedures, predictive situations, inverse problems and combination of observations. Statistics has much to offer in all these connections.

9. Update. Since the lectures were presented in 1983 work has progressed on various of the problems covered. Abrahamson (1985) has employed Smart 1 data to better see the movement of the fault rupture tip. Chiu (1986) studies the problem of estimating the parameters of a moving energy source. Lindberg (1986) develops "optimal" tapers to employ in the estimation of the frequencies of free oscillations. The book, Udias *et al* (1985) goes into substantial detail over

the formal estimation of fault plane parameters. Brillinger (1985) develops a maximum likelihood statistic for detection and estimation of a plane wave given array data. Donoho *et al* (1986) have developed a robust/resistant procedure for better aligning the seismic traces of a section. Shumway and Der (1985) indicate how the EM-method may be employed to deconvolve pulses hidden in seismic traces. The thesis, Ihaka (1985), has been completed. Ogata, (eg. Ogata (1983) and Ogata and Katsura (1986)), has carried out a variety of likelihood based analyses of earthquake times as a point process. Many statisticians have begun working on statistical aspects of inverse problems. We specifically mention O'Sullivan (1985).

IV. NEUROPHYSIOLOGY

"..., modern biometry is the interdisciplinary endeavor to build structural stochastic models of biological phenomena."

J. Neyman (1974)

10. The Field and Its Goals. Neurophysiology is the branch of science concerned with how the elements of the nervous system function and work together. The functioning is seen to involve chemical mechanisms, electrical mechanisms and physical arrangement. The studies extend from the movements of individual ions, through to the mass behavior of the components of the brain.

The goals of neurophysiologists range to the heroic: how to explain things like memory, emotion, learning, sleep, expectation, behavior. At a less ambitious level neurophysiologists are concerned with how a single nerve cell responds to stimuli, transmits information and changes with alterations of the environment.

The neuron is both the functional and structural unit of the nervous system. The brain is a multiprocessor of dramatic complexity. The elements of the nervous system may be said to differ from those in the seismic case, in that they apparently have purposes.

The field is largely experimental with researchers collecting varied and extensive data sets. The data include: photographs made via electron microscopes, fluctuating voltages and current levels within single nerve cells and finally electroencephalograms, (the brain's electrical potential at points near the skull.) The studies are sometimes simply observational, but often complex experimental designs are employed.

Important techniques that are made use of include: staining to identify individual neurons, insertion of microelectrodes to make measurements within individual cells, and the averaging of whole suites of responses to a stimulus of interest in order to reduce what can be the dominant effects of noise. Many experiments are computer controlled and computer processed.

Discoveries of neuroscientists include: nerve cells communicate with each other in both chemical and electrical fashion, the voltage pulse that travels along a neuron's output fibre is of near constant shape, there are a broad variety of nonlinear phenomena that occur. A number of verifiable physical laws and effective deterministic models (such as the Hodgkins-Huxley equations) have been set down. Much insight has been gained, especially at the level of small groups of neurons. At the level of the brain itself, knowledge is mainly phenomenological. Here the brain is viewed as a black box and studied by system identification techniques. Whatever the approach, discoveries have been made leading to lifesaving and life improving clinical diagnoses.

Statistical methods entered with the quantification of the field. No single individual scientist seems to have had a dominating effect, rather there have been many contributing workers - researchers concerned with EEG's and researchers concerned with small collections of neurons. Statistical methods entered both because of high noise levels and because a variety of phenomena seemed to be inherently stochastic. Evidence for this last is presented in Burns (1968) and Holden (1976). Pertinent books on neurophysiology include: Freeman (1975), Aidley (1978) and Segundo (1984). General reviews of statistical models and methods in neurophysiology are given in Moore *et al* (1966) for the cases of single neurons and of small groups of neurons and by Glaser and Ruchkin (1976) for EEG's. Statistical methods for classification and pattern recognition, for handling artefacts and for data summarization are in common use.

Neurobiology is one of the most active branches of science. The physiological phenomena with which it is concerned are fundamental and in most cases barely understood.

11. Neuronal Signaling. One of the important means by which nerve cells comunicate is via spike trains. (The inlays at the tops of the three graphs of Figure 13 give spike times representative of three different sorts of behavior; pacemaker (near-periodic), bursting (activity occurs in bursts), and bursting with acceleration (of firing within bursts.)

Suppose that a neuron fires at times τ_n , $n =0,\pm1,\pm2,....$ A convenient formal representation of its temporal behavior is provided by writing

$$Y(t) = \sum_n \delta(t - \tau_n)$$

with $\delta(.)$ the Dirac delta function. This representation leads to results analagous to ordinary time series results in many cases. In the case that the τ_n are random, one has a stochastic point process $\{\tau_n\}$. A principal descriptor of a point process is provided by its rate function. This is given by

$$\lim_h Prob \{point \ in \ (t,t+h]\}/h$$

as h tends to 0. In the stationary case, where the stochastic properties of the process do not depend on the time origin, the rate function is constant and so only crudely useful then.

The autointensity function is an important parameter in the stationary case. It is defined as

$$\lim_h Prob \{point \ in \ (t,t+h] \ | \ point \ at \ 0\}/h$$

as h tends to 0. This parameter may be used for example to describe the behavior of spontaneously firing neurons. Figure 13 presents examples for three cases. In the first case the neuron is firing approximately periodically. The (estimate of) the autointensity is seen to oscillate (with period equal to the interval between the points.) In the second case the neuron is evidencing activity in bursts. The probability that the neuron fires again soon after it has fired is high. In the third case the neuron is also firing in bursts; however now there is structure within the bursts, the rate of firing is seen to increase therein. The bursts here are at regular intervals.

The autointensity functions have been estimated, for this figure, by the statistic

$$\# \{ \ | \ \tau_n - \tau_m - t \ | \ < h/2\}/Nh$$

with N the total number of points and h a small binwidth. The data analysed are for the cell L10 of *Aplysia californica* , the sea hare. They were collected and

previously analysed by Bryant, Marcos and Segundo (1973). The experimental procedures and details of the data preparation may be found in that reference.

A question that arises in the study of small networks of neurons is which neurons are interacting with which? In other words, which spike trains are associated with which others? A useful parameter to employ in the study of such questions is provided by the crossintensity function. Supposing one has spike trains named M and N, then the cross-intensity function of N given M, at lag t is defined as

$$\lim_h Prob \{N \ point \ in \ (t, t+h] \ | \ M \ point \ at \ 0\}/h$$

as h tends to 0. If the M spike train consists of points σ_m and the N train of points τ_n, then this crossintensity may be estimated by

$$\# \{ \ | \ \tau_n - \sigma_m - t \ | \ < \ h/2 \}/Mh$$

with M denoting the number of M points in the data set and h a small binwidth. Figure 14 presents three examples of estimated crossintensity functions. The first graph refers to data from cells L3 and L10 of *Aplysia californica* The behavior exhibited here is that of negative association; L10's firing is inhibiting the firing of L3 (for approximately .5 second.) If one asks whether the values at negative lags differ from the level of no-association by more than sampling fluctuations, one finds they do not. This result is consistent with the cell L10 driving the cell L3. The middle graph corresponds to positive association. It is for a cell in the right viscero-pleural connective (RVP) and cell R15. The first cell tends to excite the second for about .25 seconds. The final graph represents a more complicated (polyphasic) situation. These data sets were also analysed in Bryant *et al* (1973) where further details may be found. The approximate sampling distributions of such statistics were developed in Brillinger (1975). It was found, for example, that it could be more convenient to graph the square root of the estimate in some circumstances.

The crossintensity function, being a point process analog of covariance, may be expected to be an inadequate measure of relationship. In the case of elementary statistical data it is usual to turn to regression as a better technique. In the point process case it is possible to carry out regression type analyses. For example one may fit the following form of model

$$\lim_h Prob \{N \ spike \ in \ (t, t+h] \ | \ M \ spike \ train \}/h \ = \mu + \sum_m a(t - \sigma_m)$$

as h tends to 0. The function $a(t)$ appearing in this model is referred to as the impulse response. This model may be fit as follows. Set

$$d_M^T(\lambda) = \sum_{m=1}^{M} \exp\{-i \lambda \sigma_m\}$$

with a similar definition for $d_N^T(\lambda)$. These are point process analogs of the empirical Fourier transform, (3), of time series data. The cross-periodogram of the given data at frequency λ is defined as

$$I_{NM}^T(\lambda) = (2\pi T)^{-1} d_N^T(\lambda) \overline{d_M^T(\lambda)}$$

If the cross-periodogram is smoothed to obtain $f_{NM}^T(\lambda)$, then $f_{NM}^T(\lambda)$ is an estimate of the cross-spectrum in the case that $\{M, N\}$ is a bivariate stationary point porcess. Now $A(\lambda)$, the Fourier transform of the impulse response $a(t)$, may be estimated by $f_{NM}^T(\lambda) f_{MM}^T(\lambda)^{-1}$. The impulse response itself may be estimated by back Fourier transforming A^T. The strength of the relationship proposed in the model may be measured, at frequency λ, by the sample coherency function $R_{NM}^T(\lambda) = f_{NM}^T(\lambda)/\sqrt{f_{MM}^T(\lambda) f_{NN}^T(\lambda)}$. Its modulus squared is called the sample

coherence. The coherence lies between 0 and 1, being nearer to 1 the stronger the relationship. More details of these computations may be found in Brillinger (1975) and Brillinger *et al* (1976).

Figures 15 and 16 provide the results of such an analysis for the first two data sets of Figure 14. In each case the first graph is of the sample coherence. The coherences are at some distance from the value 1.0 but above the 95% null significance level, (given by the horizontal lines in the figures.) The relationship is inherently nonlinear, so it could have been anticipated that the coherence estimate would not be close to 1.0. Further discussion of these and similar analyses may be found in the just cited work.

12. Assessing Connectivities. Questions that can arise with small networks of neurons include; is one neuron driving the rest and if one apparently is, which one is it? The next data analysis to be presented concerns three *Aplysia* cells L2, L3 and L10. From other experiments the neurophysiologists knew that cell L10 was driving cells L2 and L3. It was not known if there were any direct connections between L2 and L3. The first three graphs of Figure 17 present estimates of the three coherences, L10 with L2, L2 with L3, L10 with L3. As might have been anticipated, these suggest relationship exists in each case.

It is possible to address the question of the direct connection of cells L2 and L3, in the presence of L10, by partial coherence analysis. Suppose that $\{A,B,C\}$ is a trivariate stationary point process. Let $R_{AB}(\lambda)$ denote the coherency function of processes A and B, with similar definitions of R_{AC} and R_{BC}. Then the partial coherency of the processes B and C, having removed the (linear time invariant) effects of process A is given by

$$R_{BC \mid A} = \frac{R_{BC} - R_{BA} R_{AC}}{\sqrt{(1-\mid R_{BA} \mid^2)(1-\mid R_{CA} \mid^2)}} \tag{13}$$

suppressing the dependence on λ. This definition may be motivated several ways. One is: it is the coherency between the processes resulting when their best linear predictors based on A are removed. A second is: it is given by

$$\lim_{T \to \infty} \mid corr \{d_B^T - \frac{f_{BA}}{f_{AA}} d_A^T, d_C^T - \frac{f_{CA}}{f_{CC}} d_A^T\} \mid^2$$

Here *corr* denote the (complex) correlation coefficient. An estimate may be determined by substituting estimates for the quantities appearing on the right hand side of expression (13). If there is no connection between the processes B and C beyond their joint dependence on A, then the sample partial coherence, $R_{BC \mid A}^T$, may be expected to be near zero.

The final graph of Figure 17 provides the results of the computation for the cells L2, L3, L10 referred to. There is no suggestion of a direct connection being present. Further discussion and other examples of partial coherence computations may be found in Brillinger *et al* (1976).

13. A Structural Stochastic Model. The analyses of neuronal firing, so far presented, are of the correlation and regression type. Parameters with direct biological interpretations have not been introduced. In Brillinger and Segundo (1979) a mechanistic model is constructed and fit by the method of maximum likelihood. The model involves the following elements.

Input to a nerve cell leads to electrical current genesis. This current flows to a trigger zone, being filtered in the course of its passage. When the voltage level at the current zone exceeds a threshold value, the nerve cell fires. The neuron remembers back only to the time of previous firing. This process may be

specified analytically as follows. Let $U(t)$ denote the voltage (membrane potential) at the trigger zone at time t. Let $B(t)$ denote the time elapsed since the neuron last fired. Let $X(t)$ denote the (measured) input to the cell. Then, assuming linearity and time invariance, one can write

$$U(t) = \int_0^{B(t)} a(s)X(t-s)ds$$

for some summation function $a(.)$. The neuron fires when the process $U(t)$ crosses a threshold level $\theta(t)$.Depending on the level at which the threshold is set and the internal mechanics of the nerve cell, the input will either accelerate (excite) or slow (inhibit) the firing. In Brillinger and Segundo (1979) this mechanism was completed and discretized as follows. Input to the cell was written X_t, $t=0,...,T-1$. Corresponding output was Y_t, $t=0,...,T-1$ with $Y_t = 1$ if there was a firing in the (small) interval immediate to t and with $Y_t = 0$ otherwise. With B_t denoting the time elapsed at t since the preceding time that $Y = 1$, they set

$$U_t = \sum_{s=0}^{B_t-1} a_s X_{t-s}$$

The presence of B_t in the model had the effect of introducing a form of feedback. Finally they assumed that the threshold function had the form $\theta_t = \theta+\epsilon_t$ with the ϵ's independent standard normals.

The likelihood function of the given data and model then took the form

$$\prod_{t=0}^{T-1} \Phi(U_t-\theta)^{Y_t} \left(1-\Phi(U_t-\theta)\right)^{1-Y_t}$$

Parameter estimates were determined by maximizing this likelihood with respect to θ and the a_s. Approximate standard errors were determined by procedures traditional to maximum likelihood.

Figure 18 presents the results of one such analysis. In this case fluctuating current, $X(t)$, was injected directly into the cell R2 of *Aplysia*. The current level was taken to have marginal distribution that was approximately uniform,(but that is not crucial to the technique.) The sampling rate was 50 samples per second. The top graph of the figure gives a stretch of the noise signal inserted and the corresponding times at which the neuron fired. It is very difficult, if not impossible, to see a connection between these two stretches of data. The middle graph gives the estimated summation function, $\hat{a}_s$. The bottom graph is one means of assessing the fit of the model. It is analagous to expression (9) of Section 3, and given by

$$\# \{Y_t = 1 \ with \ U-h < \hat{U}_t < U+h\}/\# \{t \ with \ U-h < \hat{U}_t < U+h\} \qquad (14)$$

for small h, plotted versus U. Here

$$\hat{U}_t = \sum_{s=0}^{B_t-1} \hat{a}_s X_{t-s}$$

The smooth curve is the corresponding Φ. The fit may be described as adequate. The computations were carried out by a variant of the program developed for handling the seismic first motion data. Details of the experiments are given in Bryant and Segundo (1976). Further examples and discussion may be found in Brillinger and Segundo (1979). Other types of input are employed and alternate estimating procedures compared there. The large sample properties of such estimates may be studied as in Sagalovsky (1982). A great advantage of the model building approach, of this section, is that the parameters introduced and

estimated have biological interpretations. A further advantage of the maximum likelihood approach, over that of partial coherency, is that the spike trains involved can be nonstationary.

14. Analysis of Evoked Responses. A traditional means of studying the nervous system involves applying sensory stimuli to a subject and examining the ongoing electroencephalogram for an evoked response. The stimulus may be auditory, visual (eg. light flash, checkerboard pattern), olfactory, somatosensory (eg. an electrical shock), gustatory or a task. Generally the stimulus is applied for a time interval that is brief in comparison to the duration of the response. Evoked response experiments play an essential role in quantitative biology. Because the experimenter is able to chose which stimuli to apply, and when to apply them, conclusions can pass beyond associations noted to formal inferences concerning causal mechanisms.

Some dramatic success stories of the technique may be mentioned. One is presented in Bergamini *et al* (1967). Siamese twins were joined in such a way that it was not possible to determine by traditional means if the peripheral nervous pathways were interconnected. Before operating it was crucial to determine the interconnections of the twins. Ongoing EEG's were recorded for each. A series of trials were carried out in which each of the twins' legs were stimulated in turn by electrical shocks. What was found was that when a leg of one twin was stimulated, response was noted only in her EEG. On the basis of this information the twins were separated - successfully. A second notable example of the use of the evoked response technique is provided by hearing exams for newborn infants (including infants asleep.) Ongoing EEG's are recorded. These are examined for responses after loud clicks are made near the infants' ears. Rapin and Graziani (1967) present an example for an infant with hearing difficulties, both wearing and not wearing a hearing aid. The hearing aid is found to have an objectively measurable effect.

Figure 19 presents an example of evoked response data recorded at a 4 by 4 array of sensors implanted in a rabbit. In this case the stimulus was an odour and the sensors were implanted in order to study the rabbit's olfactory system. These responses were recorded concurrently. A second example is given in Figure 20. It gives the 20 successive responses evoked by applying a current pulse to the lateral olfactory tract of a rabbit and recording from a sensor implanted in the depth of the pre-piriform cortex. The signal is fairly pronounced in Figure 19. In Figure 20 the strength of the stimulus was weak and the signal is not apparent. Both of these data sets were collected in the laboratory of W. J. Freeman, University of California, Berkeley. Some details of his experiments may be found in Freeman and Schneider (1982).

Crucial to many evoked response experiments is the fact that it is generally insufficient to apply a stimulus once. Rather it must be applied repeatedly, (perhaps thousands of times), and the responses averaged. In the twins and infant examples above M equalled 250 and 100 respectively. Formally, if $Y(t)$ denotes the ongoing EEG and the stimulus is applied at times σ_m, $m = 1,...,M$, then it is usual to take as the basic statistic, the average evoked response

$$\bar{Y}(s) = \frac{1}{M} \sum_{m=1}^{M} Y(s + \sigma_m)$$

Figure 21, left-hand column, presents the results of averaging the data of Figure 20 with $M = 3, 5, 10, 20$ and 38. With increasing averaging a signal is slowly appearing from the noise. Some alternate evidence for the presence of a signal is provided by the results of the right-hand column. These are averages of 38 responses where the stimulus has been applied at a succession of increasing

strengths.

A variety of questions, that have statistical formulations, arise in the course of work with evoked responses. These include: 1. does an applied stimulus elicit a response? 2. do two different stimuli elicit the same response?, 3. is the same response elicited at two different sensor locations?, 4. is the response stationary?, 5. if the order of stimuli application is altered, are the corresponding responses altered?, 6. are the efects of different stimuli additive?, 7. how does the response depend on the stimulus intensity?, and 8. how do the responses depend on exogenous variables? To go with answers to these questions, researchers seek quick efficient data collection, precise estimates and indications of variability. Difficulties that commonly arise include: small response, large noise, variability in response, artifacts present and effects are superposed. Next in this section, two formal set-ups will be presented that may be employed to address the situation.

Suppose, to begin, that there is a single stimulus and that it is applied at times σ_m. Let $a(.)$ denote the response in a single shock experiment. If the system is time invariant, and the effects of the various shocks additive (superposable), then a model for consideration is

$$Y(t) = \mu + \sum_m a(t - \sigma_m) + \epsilon(t) \tag{15}$$

with $Y(.)$ denoting the ongoing EEG and $\epsilon(.)$ denoting noise. In the case of the EEG this model seems to have to be empirically verified, rather than being an implication of basic biology. For example, the assumption of superposability may be examined as follows. To begin, carry out some single shock experiments, i.e. apply the shocks at times far enough apart that their individual effects seem likely to have died off. Let $\hat{a}(s)$ denote the average of the responses evoked, with s lag since stimulus application. Now carry out some two shock experiments, i.e. apply shocks say Δ time units apart. Let $\hat{b}(s,\Delta)$ denote the average of the responses evoked. To examine the assumption of superposability compare $\hat{a}(s) + \hat{a}(s-\Delta)$ with $\hat{b}(s,\Delta)$. The results of carrying out such a check, in an experimental situation, are given in Biedenbach and Freeman (1965). They form averages of $M = 150$ responses, and do not note departure from superposability.

We now turn to one formal analysis of the model (15). If one writes

$$X(t) = \sum_m \delta(t - \sigma_m)$$

then (15) takes the form

$$Y(t) = \mu + \int a(s)X(t-s)ds + \epsilon(t)$$

i.e. it is seen to be the model of cross-spectral analysis. Taking Fourier transforms, one has

$$d_Y^T(\lambda) \approx A(\lambda)d_X^T(\lambda) + d_\epsilon^T(\lambda)$$

for $\lambda > 0$, with $A(\lambda)$ denoting the Fourier transform of $a(.)$. Consider a number of frequencies $\lambda_k = 2\pi k/T$ near λ. Then, assuming $A(.)$ smooth, one has the approximate linear model

$$Y_k \approx A(\lambda)X_k + E_k$$

with

$$Y_k = \sum_{t=0}^{T-1} Y(t) exp\{-i\frac{2\pi kt}{T}\}$$

and similar definitions of X_k, E_k. Next, via a central limit theorem for empirical Fourier transforms, the noise variates, E_k, may be approximated by independent

(complex) normals having mean 0 and variance $2\pi T f_{\epsilon\epsilon}(\lambda)$. All the inference procedures for the linear model become available. For example, as an estimate of $A(\lambda)$ one has

$$\hat{A}(\lambda) = \sum_k Y_k \bar{X}_k / \sum_k X_k \bar{X}_k$$

and this variate will be approximately distributed as complex normal with mean $A(\lambda)$ and variance $2\pi T f_{\epsilon\epsilon}(\lambda)/\sum_k |X_k|^2$. This formulation has a variety of convenient properties. It directly extends to the cases of multiple stimuli and multiple responses. It handles stimuli of varying intensity. It allows the individual responses of the separate shocks to overlap. Formal inference procedures, such as tests, are available. Complex experiments may be designed and analyzed - complexities handled such as: blocking, rotation, factorial treatment structure, measured covariates. Formal checks for interaction are available. Finally, one can turn to the question of optimal design.

It is sometimes convenient to adopt a different viewpoint for the problem. Suppose that the shocks are applied at times such that $\sigma_{m+1} - \sigma_m > V$ with $a(s) = 0$ for $s > V$ and $s < 0$. Write

$$Y_m(s) = Y(s + \sigma_m)$$

Then $Y_m(s) = \mu + a(s) + \epsilon_m(s)$ for $0 \le s \le V$. The average evoked response is now conveniently denoted $\bar{Y}(u)$. As an example of the use of this formulation, suppose there are I different stimuli and that each are applied J times, then one is led to set down the model

$$Y_{ij}(s) = \mu_{ij} + a(s) + b_i(s) + \epsilon_{ij}(s)$$

with i indexing stimuli and j indexing replicates. Other methodologies such as growth curves and discriminant analysis are seen to become available with this formulation.

It was mentioned that evoked response data may be contaminated by artefacts. It is perhaps worth noting that robust/resistant estimates are directly available. Suppose one has a measure of distance, such as

$$\| Y - a \|^2 = \int_0^V |Y(s) - a(s)|^2 ds$$

and an estimate of scale, $\hat{\rho}$. Then a family of robust/resistant estimates is provided by

$$\hat{a}(s) = \sum_m W_m Y_m(s) / \sum_m W_m$$

with $W_m = W(\| Y_m - \hat{a} \| / \hat{\rho})$ and $W(.)$ a univariate set of multipliers for robust/resistance. The estimate will need to be computed recursively. An elementary example is provided by the "trimmed mean"

$$\hat{a}(s) = \sum{}' Y_m(s) / \beta M$$

with $\sum'$ over the βM smallest $\| Y_m - \hat{a} \|$. This class of estimates was proposed in Brillinger (1981) and investigated in Folledo (1983). The top graph of Figure 22 provides an example of this estimate with 50% trimming ($\beta = .5$), in the case of data like that of Figure 20, (but with a stimulus of strength 122% of the threshold stimulus.) The sold curve denotes the average evoked response, the dashed one the trimmed statistic. The two curves are near identical, although when examined the individual responses are found to differ by a fair amount.

It is to be noted that a "real-time" version of such a trimmed mean may be computed, see Brillinger (1981). This statistic is given recursively for $m = 1,2,...$

by

$$\hat{\rho}_{m+1} = \hat{\rho}_m - \frac{L}{m}(\frac{1}{\beta}-1)$$

if $\| Y_{m+1}-\hat{a}_m \| \leq \hat{\rho}_m$ and

$$\hat{\rho}_{m+1} = \hat{\rho}_m + L/m$$

otherwise, and by

$$\hat{a}_{m+1}(s) = \hat{a}_m(s) + \frac{1}{\beta m}(Y_{m+1}(s)-\hat{a}_m(s))$$

if $\| Y_{m+1}-\hat{a}_m \| \leq \hat{\rho}_m$ and

$$\hat{a}_{m+1}(s) = \hat{a}_m(s)$$

otherwise. In preparing a worked example, it was found more convenient in the choice of L to replace ρ by its logarithm. The bottom graph of the figure gives the result for the same data as that of the top graph. The alogorithm was run setting $\hat{a}_1(s) = Y_1(s)$ and $L = .15$. The real-time estimate, given by the solid line, has performed virtually as well as the dead-time estimate in this case.

15. A Confirmed (Fourier) Inference.

Muscle cells are electrochemical devices. If the chemical acetylcholine is applied at the neuro-muscular junction, measurable voltage fluctuations result. Specifically, acetylcholine release causes postsynaptic membrane channels to open leading to voltage fluctuations. Katz and Miledi (1971,1972) measured voltage fluctuations associated with this phenomenon and found that the power spectrum could be approximated by the functional form $\alpha/(\beta^2+\lambda^2)$. (An example of the fit of this function to such data and a description of a fitting procedure may be found in Bevan *et al* (1979). They proposed the model

$$Y(t) = \sum_m a(t-\sigma_m)$$

with the σ_m points of a Poisson process and with $a(t) = \exp\{-\beta t\}$. This $a(.)$ function corresponds to the effectiveness of an open channel decaying exponentially and leads to a power spectrum of the indicated form. Katz and Miledi mentioned that the pulses might actually be rectangular of random duration, but they preferred to deal with the exponential form. Stevens (1972) proposed the specific model

$$Y(t) = \sum_m a_m(t-\sigma_m)$$

also with $\{\sigma_m\}$ Poisson, but now with $a(t) = 1$ for $0 < t < T_m$ and $a(t) = 0$ otherwise. The T_m are independent exponentials of mean $1/\beta$ and correspond to the lengths of time that the channels are open. Stephens noted that this model also led to a power spectrum of the form $\alpha/(\beta^2+\lambda^2)$. The models were indistinguishable with the data collected.

The problem was later resolved by improved experimental technique. Neher and Sakman (1976) developed a technique that allowed the opening and closings of individual channels to be seen. They found that the channels remained equally effective and open for time periods of varying lengths. The two proposed models could be distinguished.

Examples of single-channel data and the corresponding estimated power spectrum may be found in Lecar (1981). Jackson and Lecar (1979) present results confirming the exponential duration of the openings.

16. Other Topics. Spatial temporal data are commonly collected by neuroscientists. One form is the electroencephalogram recored by an array of sensors on the scalp. Figure 19 presented an example of data collected for the olfactory system of the rabbit. The stimulus was release of the odour ethylacetate. An 8 by 8 array of electrodes was imbedded in the animal. The data, already presented in Figure 19, gives the responses for the sensors at the postitions with x-coordinates 2, 4, 6, 8 and y-coordinates 1, 3, 5, 7 of Figure 23. One procedure that W. J. Freeman has found helpful for understanding this type of data is the computing of empirical orthogonal functions, see Freeman (1980). Figure 23 gives an example. These results are derived by stacking the responses into a matrix, **X**, with rows corresponding to sensor and columns to time, and then computing the singular value decomposition , $\mathbf{X} = \mathbf{UDV}'$, of that matrix. The **U** for a particular component, say the first, are then plotted versus sensor location as in the top graph of Figure 23. The **V** values are similarly plotted versus time and appear in the bottom graph. The results of Figure 23 are based on 64 series, not just the 16 of Figure 19. The contour plot suggests the presence of a focus of activity. The time series component elicited may be seen lurking in the individual responses of Figure 19. (It may be mentioned that meteorologists have long computed empirical orthogonal functions for spatial-temporal data and used them in forecasting, see Lorentz (1956) for example.)

Childers has also made use of array data in studying the neural system. In Childers (1977) he estimates the frequency-wavenumber spectrum for responses evoked by visual stimuli (light flashes) in the human EEG. He was concerned to estimate the speed and direction of propagating waves. In the paper cited he first notes an apparent high-velocity wave. After this wave has been "removed", he notes the presence of a pair of waves moving in opposite directions. His research is directed at developing a diagnostic procedure for various visual disorders and at obtaining insight concerning how the visual system functions.

The decaying cosine model of Section 2 has also found a use in neurophysiology. In his work with the olfactory system Freeman (1972, 75, 79) found that the average evoked response could be well-fitted by the sum of a few decaying cosine terms. He developed a model involving spike to wave conversion, involving collections of constant coefficient second-order differential equations, involving feedforward and feedback and involving wave to spike conversion. He employed non-linear regression in the time domain to estimate the unknowns. In one case, involving two cosines, he was led to view the stronger wave as representing intracortical negative feedback and the weaker as representing a second feedback loop. Of interest in this type of work is what happens to the frequencies and the decay rates when the experimental conditions are altered. A second reference to decaying cosines is Childers and Pao (1972). They consider the model

$$Y(t) = \sum_k \alpha_k \, t \, \exp\{-\beta_k \, t\} cos(\gamma_k \, t + \delta_k) + \epsilon(t)$$

for visual evoked responses monitored over the occipital region. In particular they study the data by complex demodulation.

Brief reference will be made to several other topics. Dumermuth *et al* (1971) estimate the bispectrum of human EEG's. de Weerd and Kap (1981) discuss the computation of some time-varying quantities. Marmarelis and Naka (1974) consider the case of biological systems with several inputs. An extreme case of this occurs when the input is varying in both time and space. This circumstance is considered in Yasui *et al* (1979). The book Marmarelis and Marmarelis (1978) goes into great detail concerning the identification of systems that are polynomial and time invariant in the input. They emphasize the advantages resulting from employing a Gaussian white noise input. The dedication of the book is worth

mentioning - "*To an ambitious new breed: SYSTEMS PHYSIOLOGISTS*"

Another area of research activity has been that of control. The works Poggio and Reichardt (1981) and Wehrhahn *et al* (1982) may be noted. They are concenerned with data that are three dimensional trajectories.

17. Discussion. As the examples presented hopefully indicate, a broad range of data types arise in the neurosciences. Further, that data is collected at both the micro level and the mass level. Procedures developed have the opportunity to move on to clinical use.

The field of neurophysiology has the satisfying aspect that often controlled laboratory experiments are possible and repeatable. Further there are opportunities for design of experiments. In the field statistics has been seen to provide techniques for model formation and validation, for measuring uncertainty in conclusions and for addressing questions of causality. Statistical techniques have led to insight concerning the underlying physiology.

It is interesting to note the evolution of the analysis in the case of the neuronal signaling analysis as presented in Sections 11 and 13. One can recognize the stages of: 1. (feature) description, 2. correlation/association, 3. (*adhoc*) regression, 4. mechanistic model. These stages are usual in many elementary situations.

18. Update. The analysis of single ion-channel data, briefly referred to in Section 15, has become a whole industry. Models with several states are now routinely fit. References include: Colquhoun and Hawkes (1983), Labarca *et al* (1986). Extending the work of Section 13 Brillinger (1986) presents a number of examples of the maximum likelihood fitting of a neural model employing corresponding spike train input and output data. Smith and Chen (1986) study a more complicated neural model. The chirp signal was propounded as being of substantial importance in seismic exploration. Some use of it has been made recently in physiological studies. In Norcia and Tyler (1985) a 10 second spatial freqency sweep stimulus is employed and the corresponding visual evoked potential measured. Finally we note that Grajski *et al* (1986) apply modern classification procedures to study the effects of applying different odours on the olfactory bulb EEG's of rabbits.

V. CONCLUDING REMARKS

In this article we have presented a number of examples, drawn mainly from our personal experience, showing the use of the same statistical technique in the rather separate sciences of seismology and neurophysiology. It now seems appropriate to ask what, if anything, have the *three* sciences - statistics, seismology, neurophysiology - gained from each other as a result of connections albethey indirect? Having in mind a broader class of examples then those discussed in this paper, one can say that: i) statistics is richer for having been led to develop and study various novel methods to handle specific problems arising in seismology or neurophysiology, ii) both seismology and neurophysiology are the richer for the other's field having generated a problem for the statistician to abstract sufficiently that the result's applicability to their field became apparent, iii) either seismology or neurophysiology benefit from a statistical formulation because various of their problems seem necessarily to need to be stated in terms of probabilities (eg. neither neuron firings nor earthquakes seem deterministic) and because these fields need procedures to validate results and to fit conceptual models. The methods of statistics can lead to important insight and understanding in

substantive problems.

It may be remarked that the applicability of statistical procedures to these two substantive fields has further grown in direct consequence of their move to greater quantification and digital data collection. The data sets analysed were of high quality. The fact that the analyses were informative to an extent here bodes well for the use of such techniques in fields with data of lesser quality.

The reader will have noted that some of the analyses were time-side and some were frequency-side. Each domain has its advantages. It is possibly worth pointing out specifically that stationarity was not required for some of the frequency-side procedures.

On review it may be seen that the techniques employed for time series data and for point process data in many cases are not that different. Brillinger (1978) presents some comparative discussion of the techniques for the two cases. Our presentation is somewhat remiss in the seismological case in not presenting some worked examples of auto and cross- intensity estimation.

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LEGENDS

Figure 1. Trace of the Chilean great earthquake of 22 May 1960 as recorded by the tiltmeter in the Grotta Gigante at Trieste.

Figure 2. The logarithm of the periodogram of the data of Figure 1 using 2548 data values.

Figure 3. The result of complex demodulating the data of Figure 1 at the frequency of 3.89 cycles per hour. The top graph gives the logarithm of the running amplitude. The bottom graph gives the running phase. The bandwidth of the filter employed is .594 cycles/hour.

Figure 4. As for Figure 3, but at the frequency of 5.67 cycles per hour.

Figure 5. The CAL 8 earth model. The scales for the velocities are km/sec. The scale for the density is g/cm^3.

Figure 6. The P-wave first motion data for the earthquake of the Alaskan sequence that took place 30 March 1964 at 0200. The solid circles refer to compressions. The open circles to dilations.

Figure 7. A plot of the statistic (9) of Section 3 for the data of 16 events of the Alaskan sequence.

Figure 8. The top graph gives the transverse shear wave component of the signal of the 29 January 1981 magnitude 6.7 Taiwan earthquake as recorded by the accelerometers of the Smart 1 array some 30 km northwest of the epicenter. The lower graph plots the difference between the amplitude of the Fourier transform of the data and a strongly smoothed version of those values versus the strongly smoothed values.

Figure 9. The amplitudes of the Fourier transforms of the Taiwan data of Figure 8 and the corresponding expected values as computed for the model of Section 4. The data stretch consisted of 256 points. Both scales of the plot are logarithmic.

Figure 10. The vertical P-wave portion of the 29 January earthquake as recorded at 9 of the sensors of the Smart 1 array. The "*" are plotted at the physical locations of the sensors.

Figure 11. The frequency-wavenumber periodogram of the data of Figure 10. The time series stretches contained 720 points. The temporal frequency was 1.94 cycles/second.

Figure 12. The high resolution (or Capon) spectrum computed for S-wave data recorded, following a strip mining blast, at the Large Aperture Seismic Array located in Montana. The temporal frequency was 1.25 cycles/second.

Figure 13. Point process data from the cell L10 of *Aplysia californica* behaving in three different fashions. The inlays at the top of the three graphs give brief stretches of the data. The functions plotted are estimates of the autointensity functions based on 1538, 1019, 1631 spikes respectively.

Figure 14. Estimates of the crossintensity functions for three pairs of *Aplysia* neurons. The estimates are based on (1746,302), (1101,288), (1019,993) spikes in the pairs of trains respectively.

Figure 15. The estimated coherence and impulse response for the data of the top graph of Figure 14. The horizontal line gives an estimate of the level exceeded by chance only 5% of the time when the spike trains are independent.

Figure 16. The estimated coherence and impulse response for the data of the middle graph of Figure 14. The horizontal line in the top graph give the upper 95% null point of the distribution of the sample coherence.

Figure 17. Data for a network of three *Aplysia* neurons. The partial coherence estimate is based on expression (13) of Section 12. In each case the horizontal solid line gives the upper 95% null level.

Figure 18. The results of fitting the neuron model of Section 13 to data obtained in an experiment with the cell R2 of *Aplysia*. The upper graph is a segment of the data. Noise, (lower trace), is introduced into the cell. The upper trace gives corresponding observed firing times. The interval between the sampling points is .02 second. The middle graph gives the maximum likelihood estimate of the summation function, $a(.)$. The bottom graph provides the statistic (14) of Section 13.

Figure 19. The bursts of activity recorded at the 16 sensors of a 4 by 4 array following the stimulation of a rabbit by an odour. The units of the x-axis are seconds.

Figure 20. Twenty successive responses evoked in the pre-piriform cortex by (electrically) stimulating a rabbit. The x-axis units are seconds.

Figure 21. The various graphs here are meant to show the effects of changing the number of responses averaged (left column) and the strength of stimulus applied (right column) for data such as that of Figure 20.

Figure 22. Top graph compares the average evoked response with the 50% trimmed mean for the data taken at 122% of a threshold value. The lower graph contrasts the 50% trimmed mean statistic with a value computed recursively.

Figure 23. The results of a singular value decomposition of the full set of the data from which the bursts of Figure 19 were taken. The values graphed are for the first components. The axes in the top graph give spatial location.

1960 Chilean Earthquake - Tides Removed

time elapsed (hours)

time elapsed (hours)
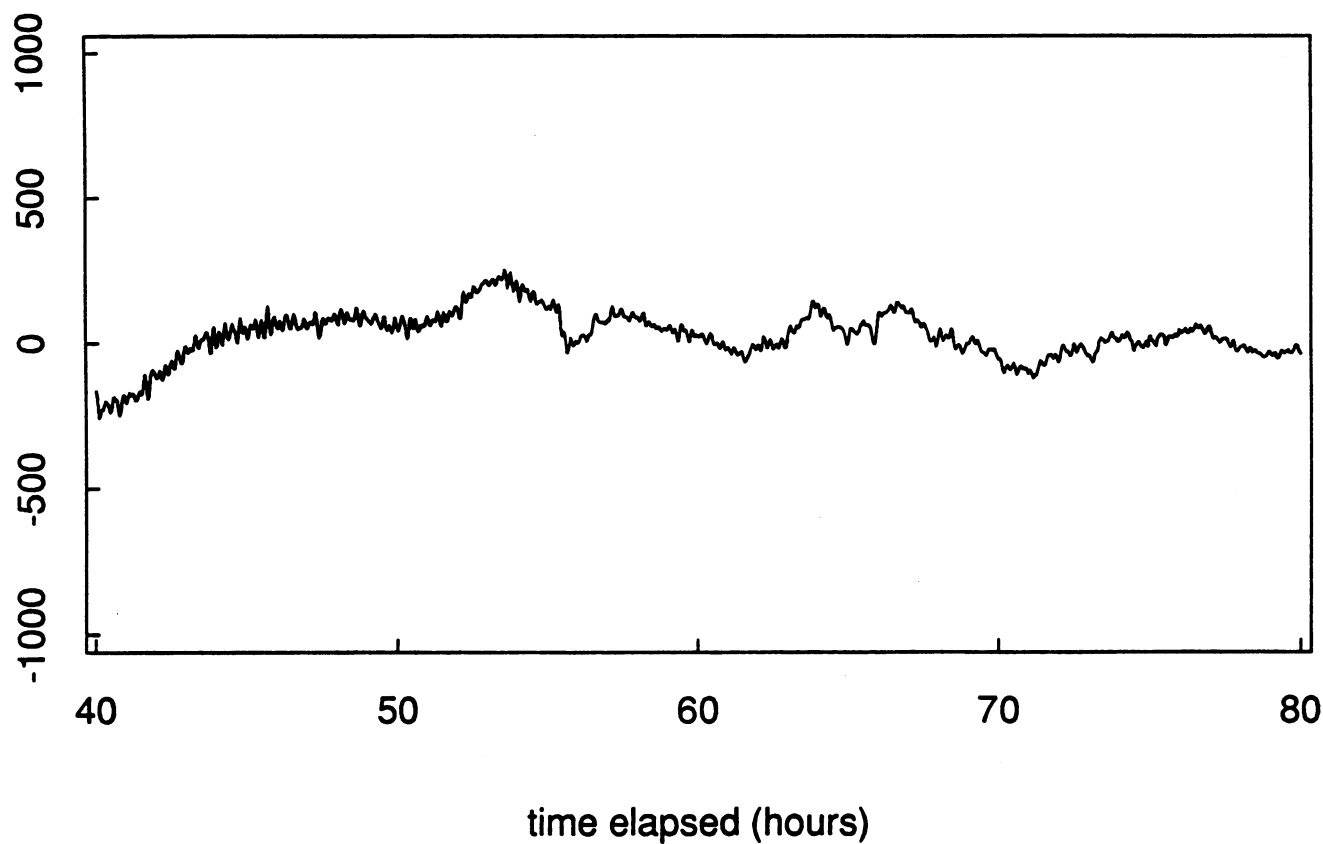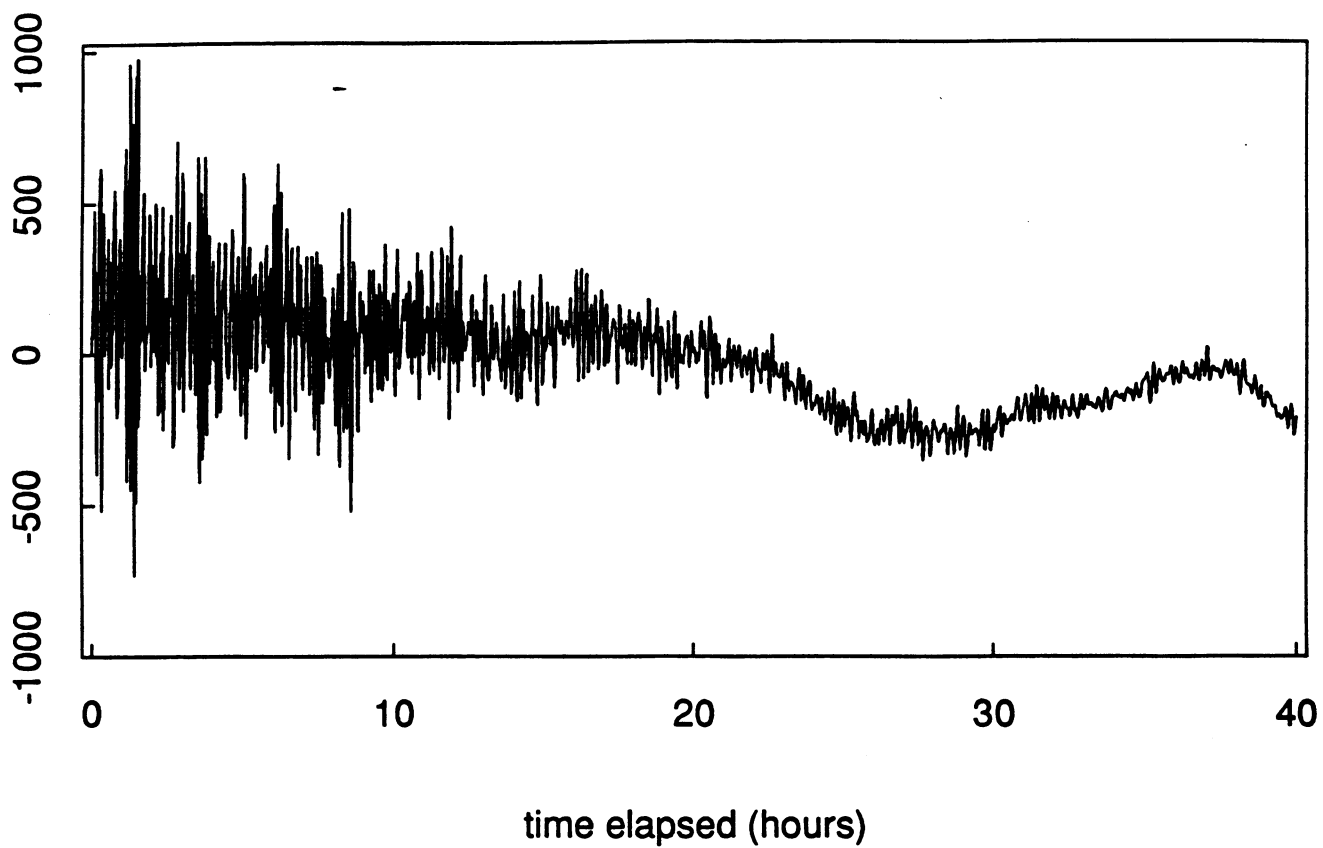
Periodogram - Chilean Data

frequency (cycles/hour)

frequency (cycles/hour)
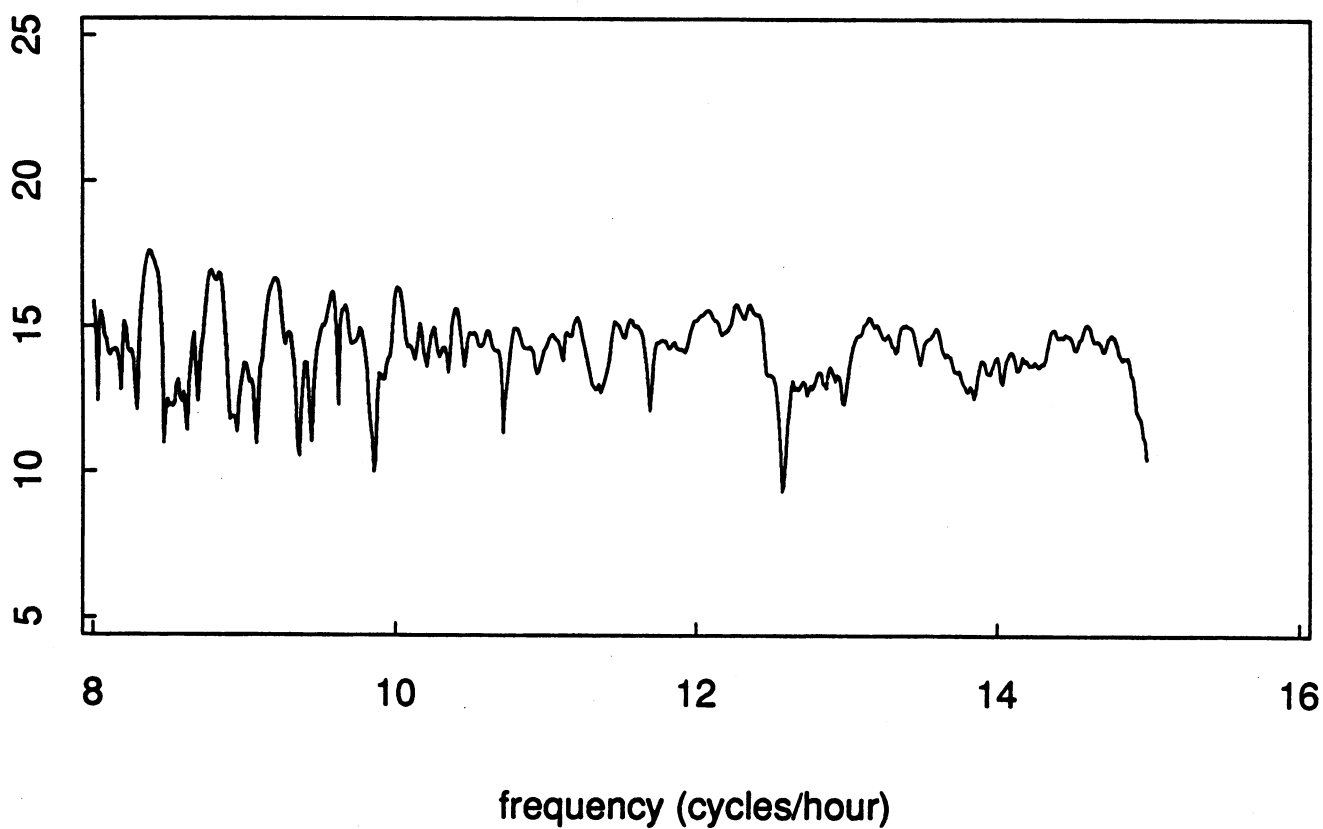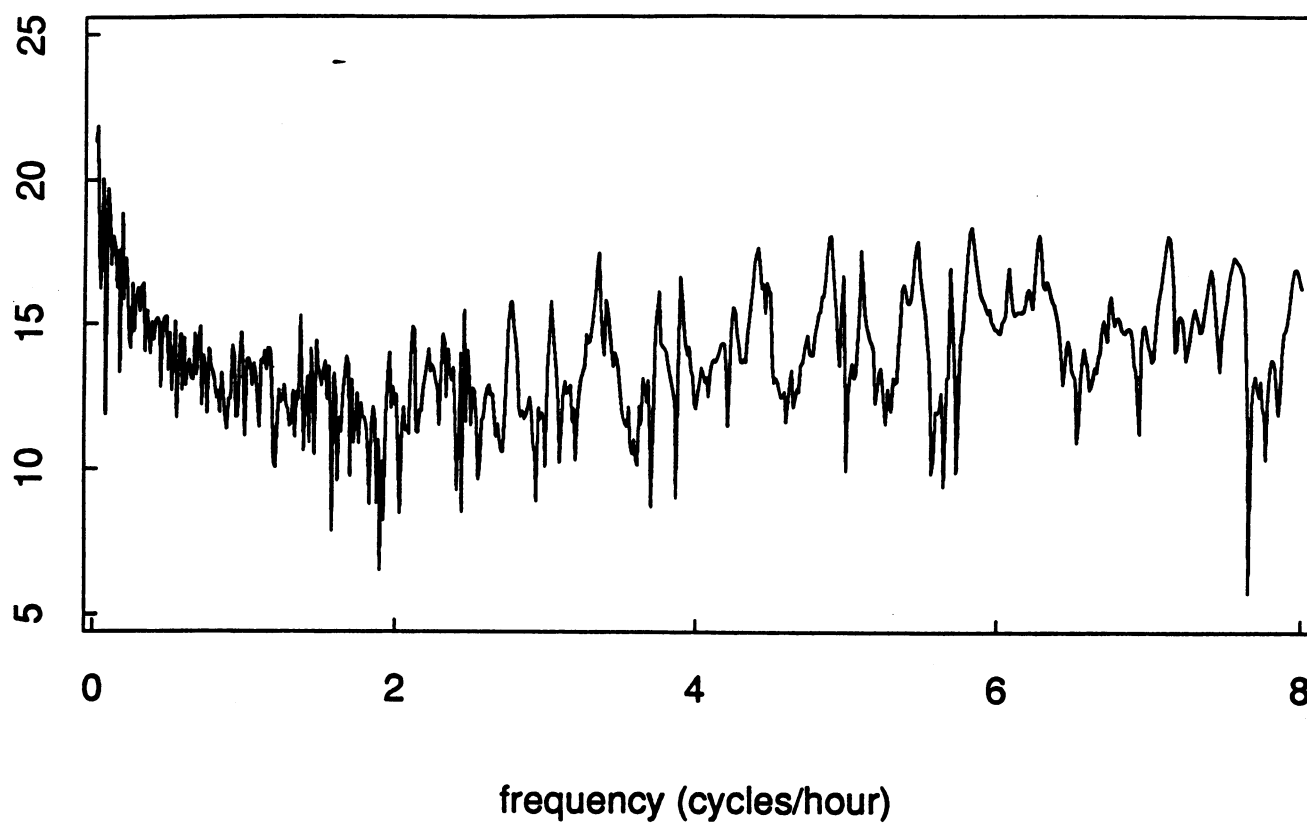
Log Amplitude at 3.89 cycles/hour

time elapsed (hours)

Phase Angle

time elapsed (hours)
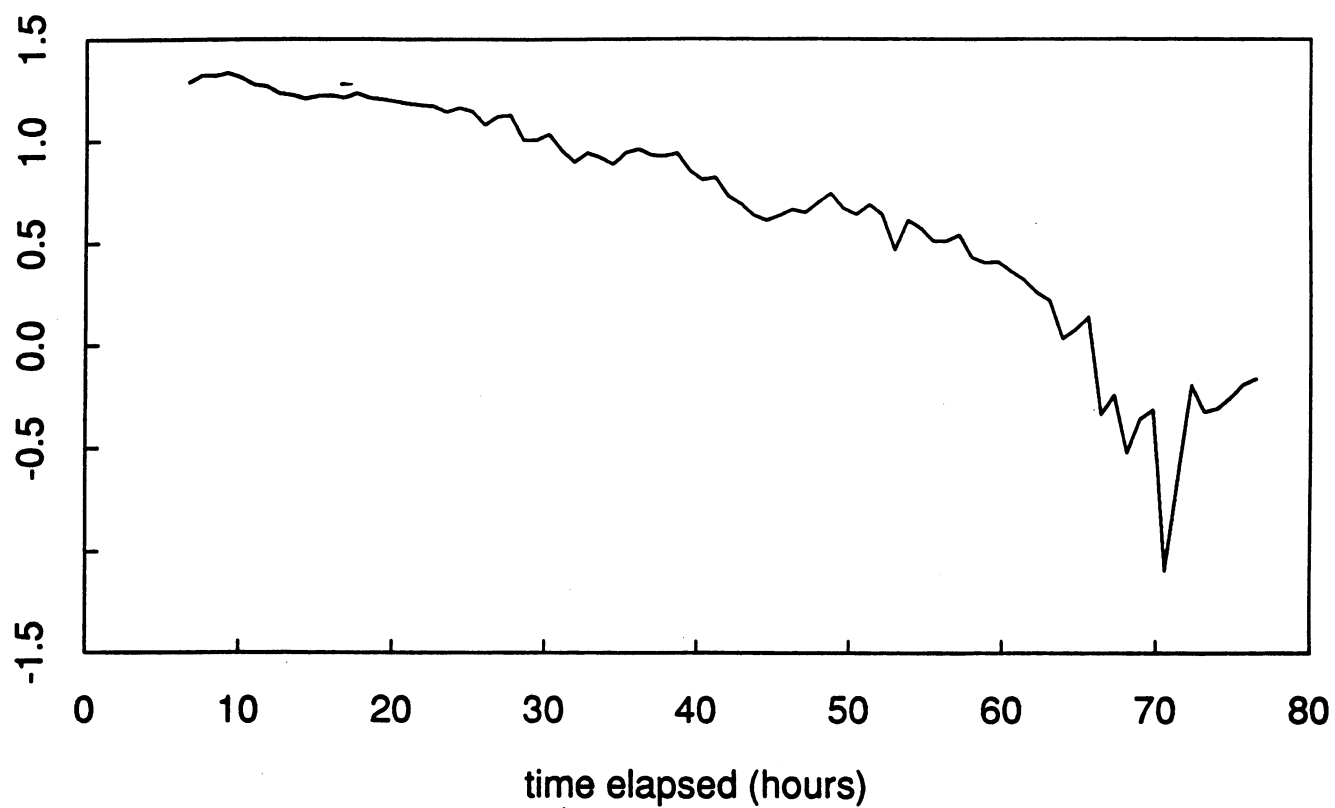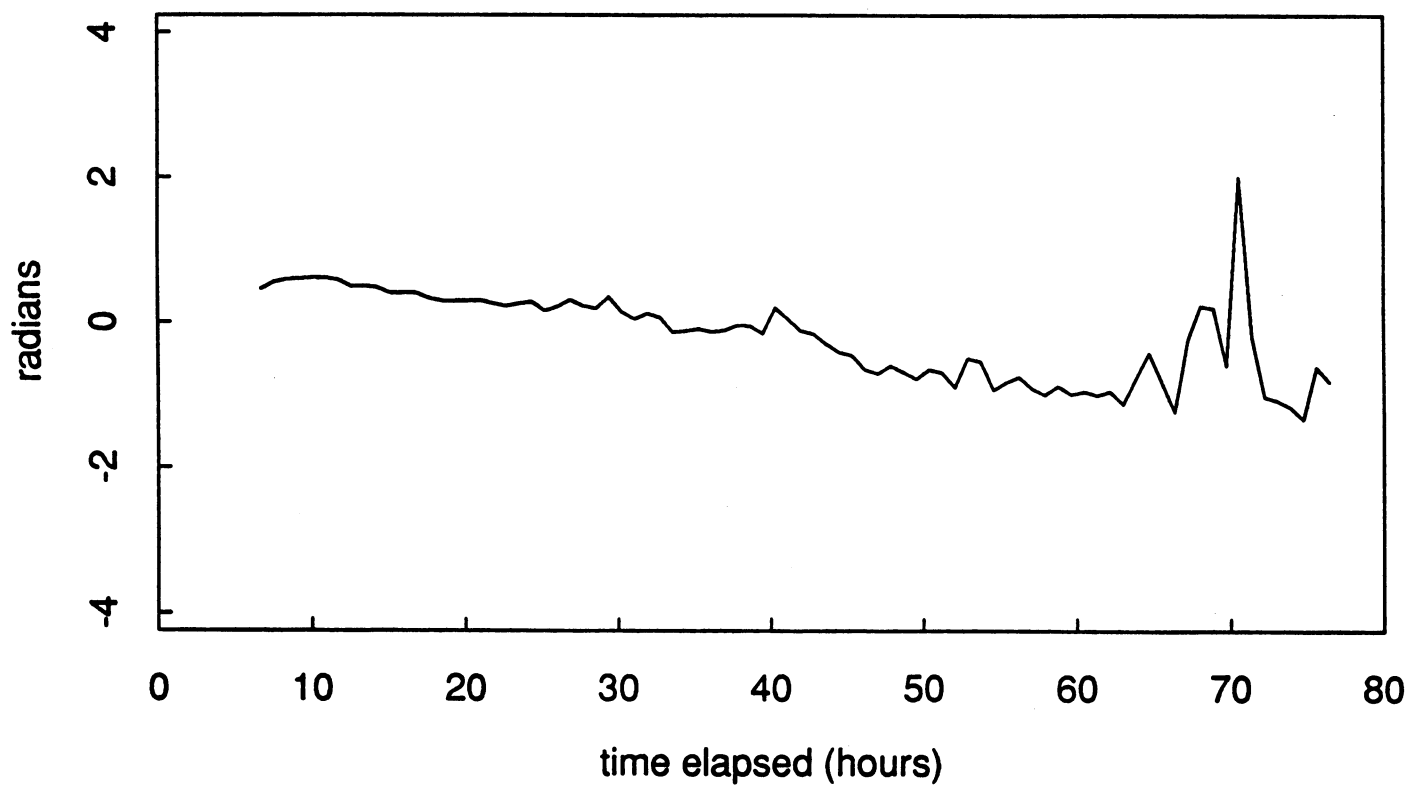
Log Amplitude at 5.67 cycles/hour

time elapsed (hours)

Phase Angle

time elapsed (hours)

Earth Model CAL8

density

S-wave velocity

P-wave velocity

depth (km)
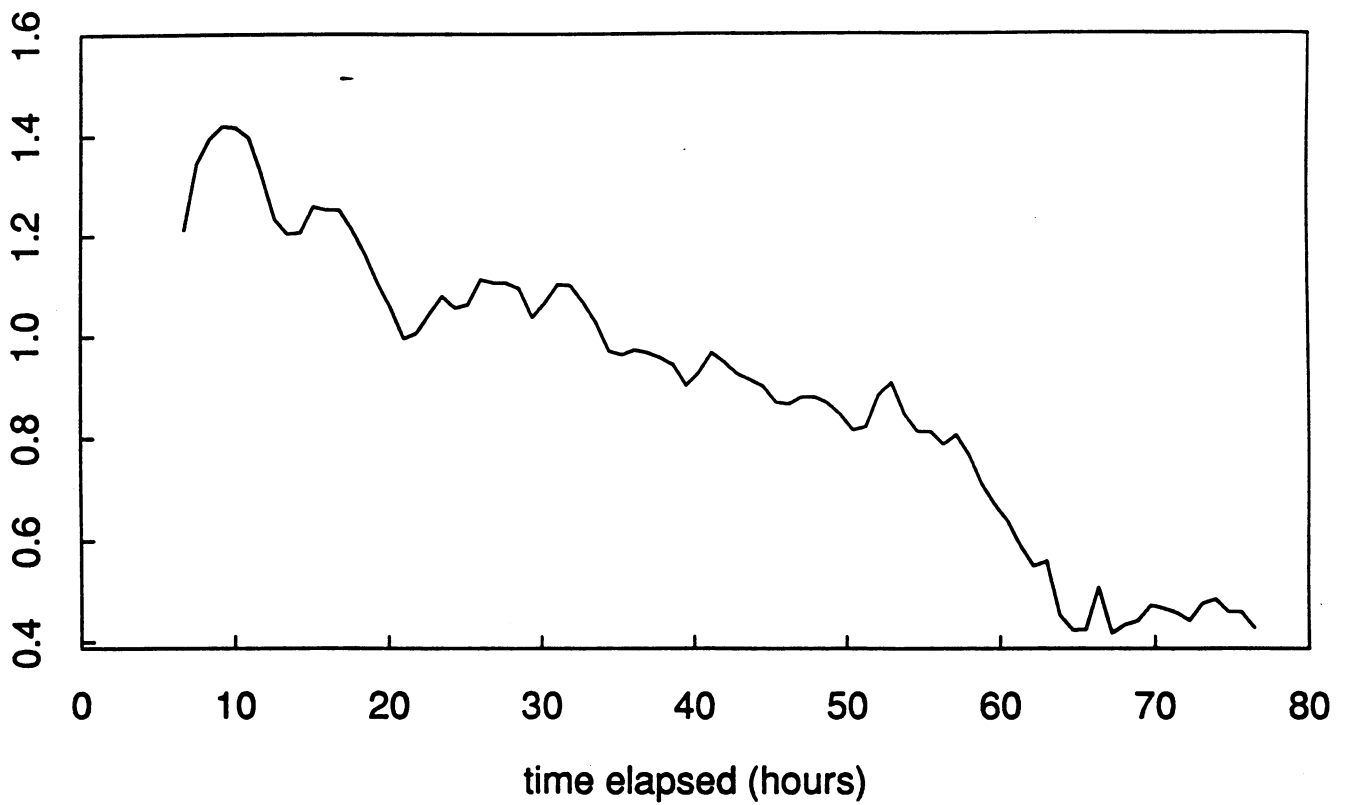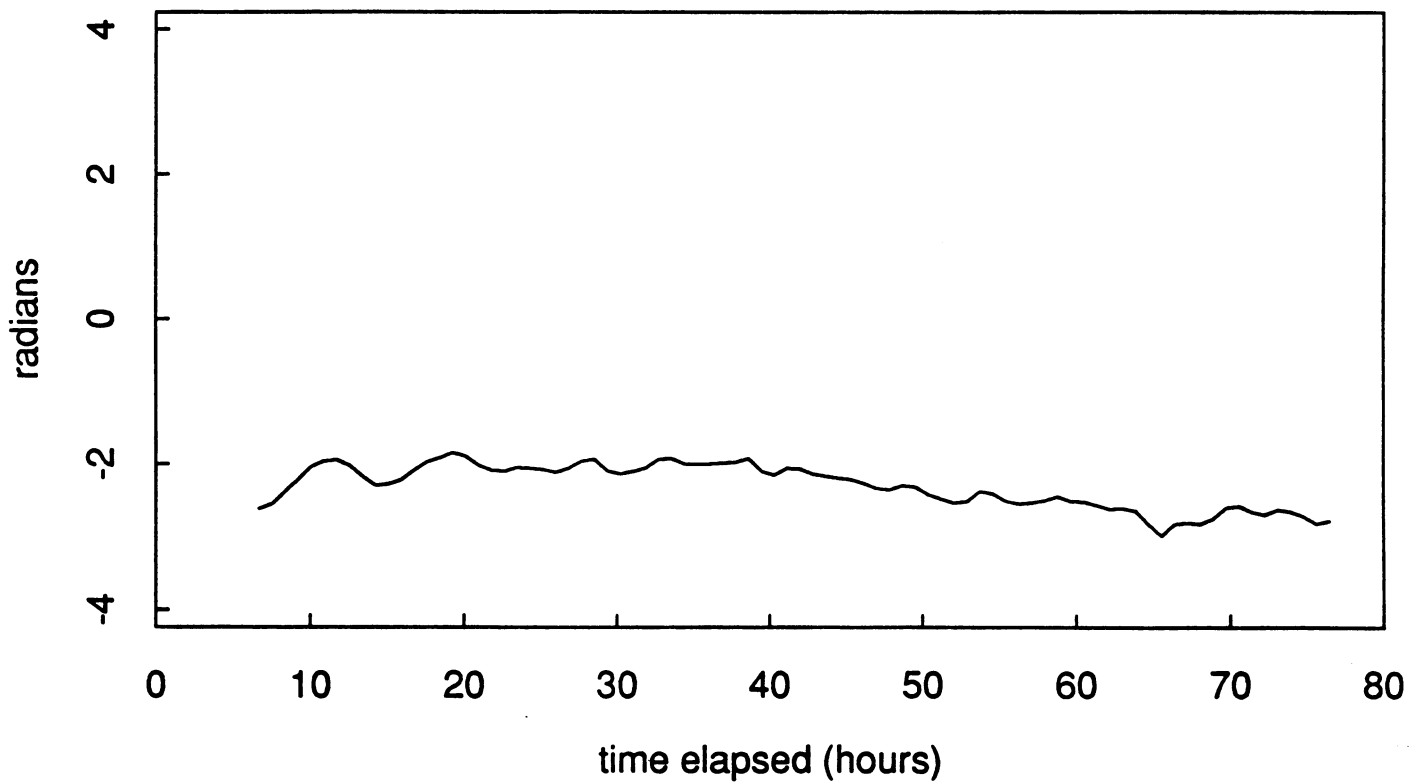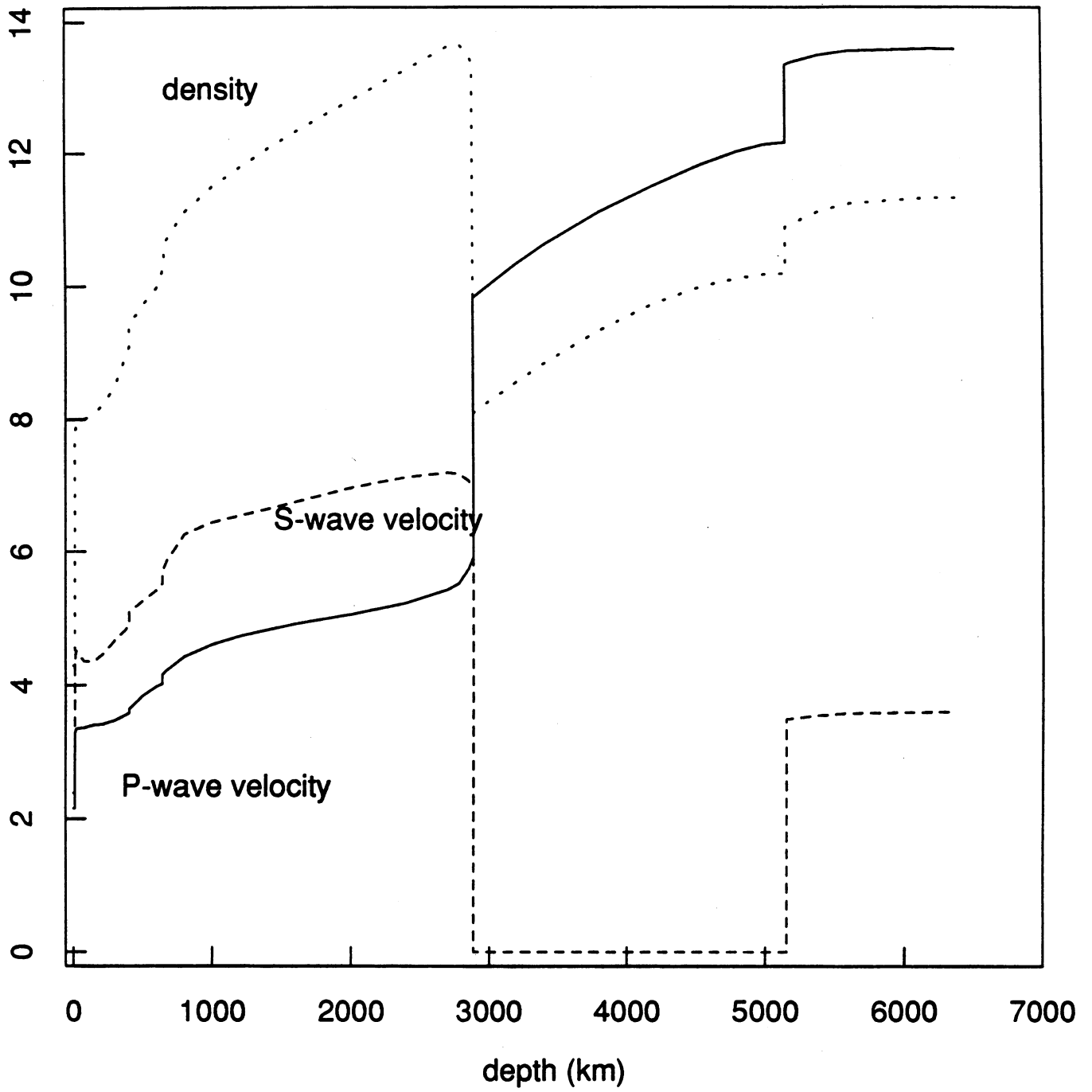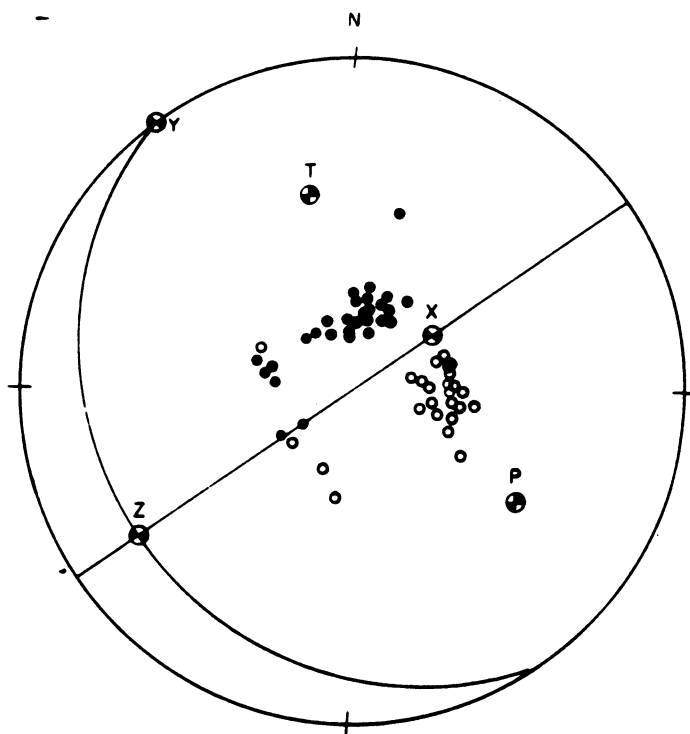
Empirical Probability of Correct First Motion

Taiwan Event - Shear Wave

seconds

Error Plot

errors

smoothed amplitude
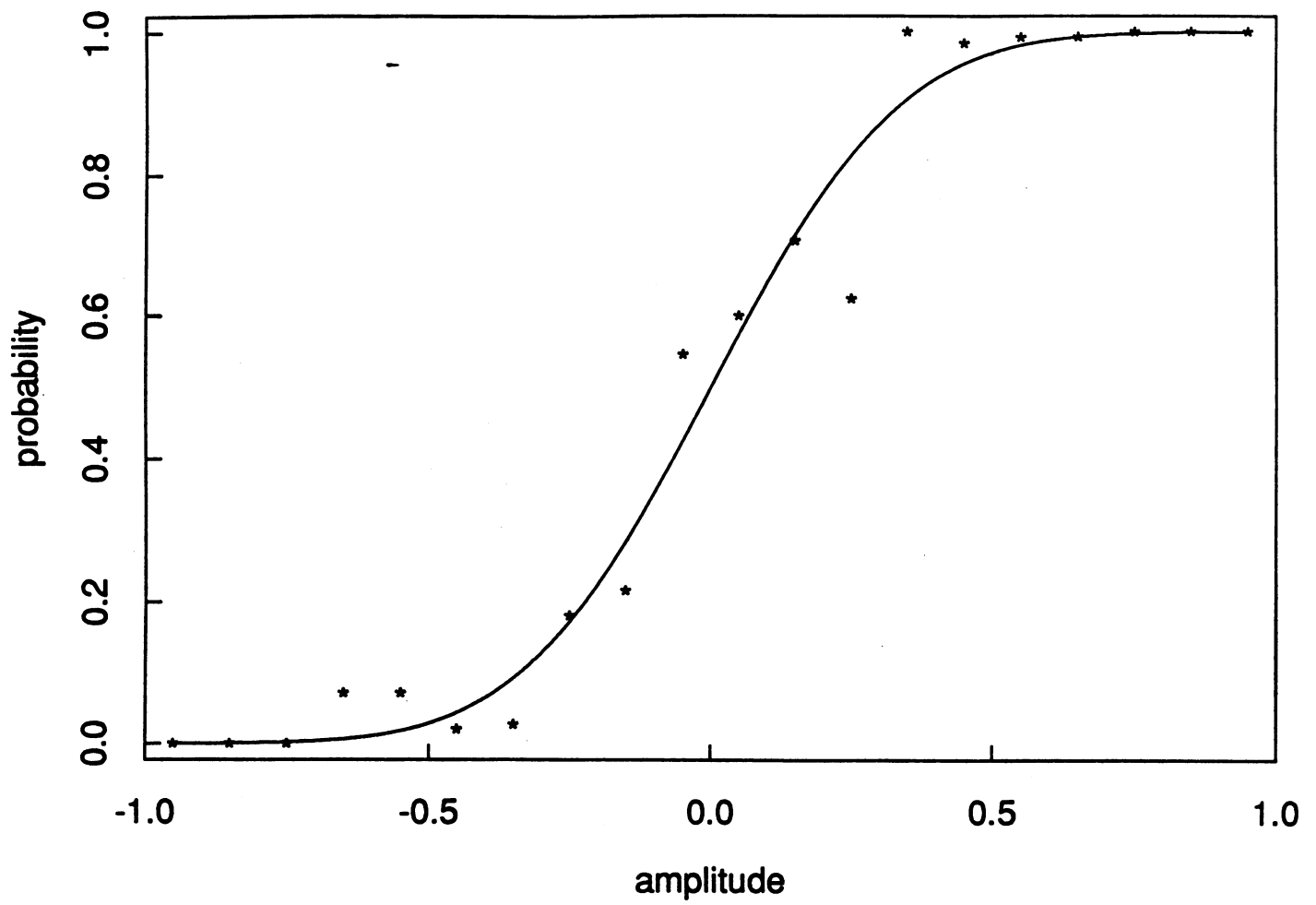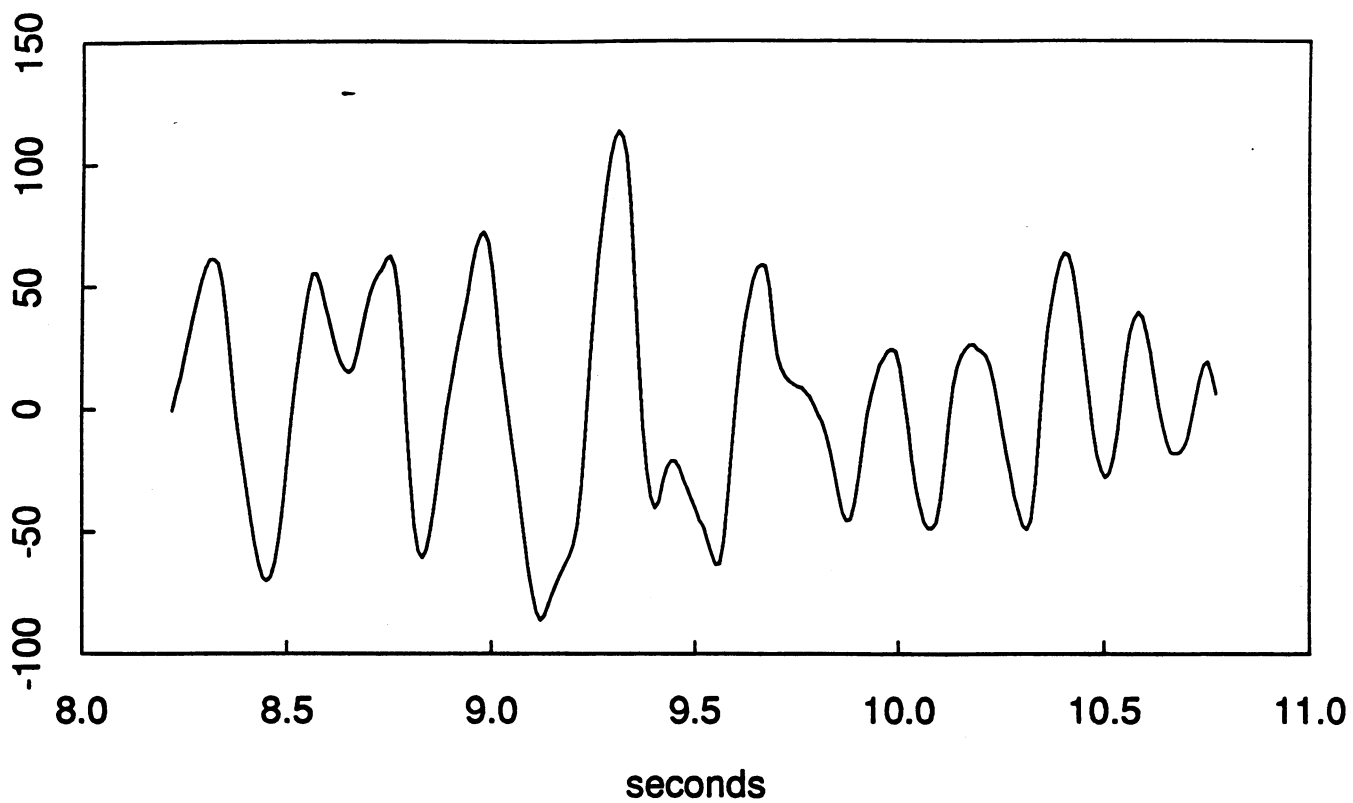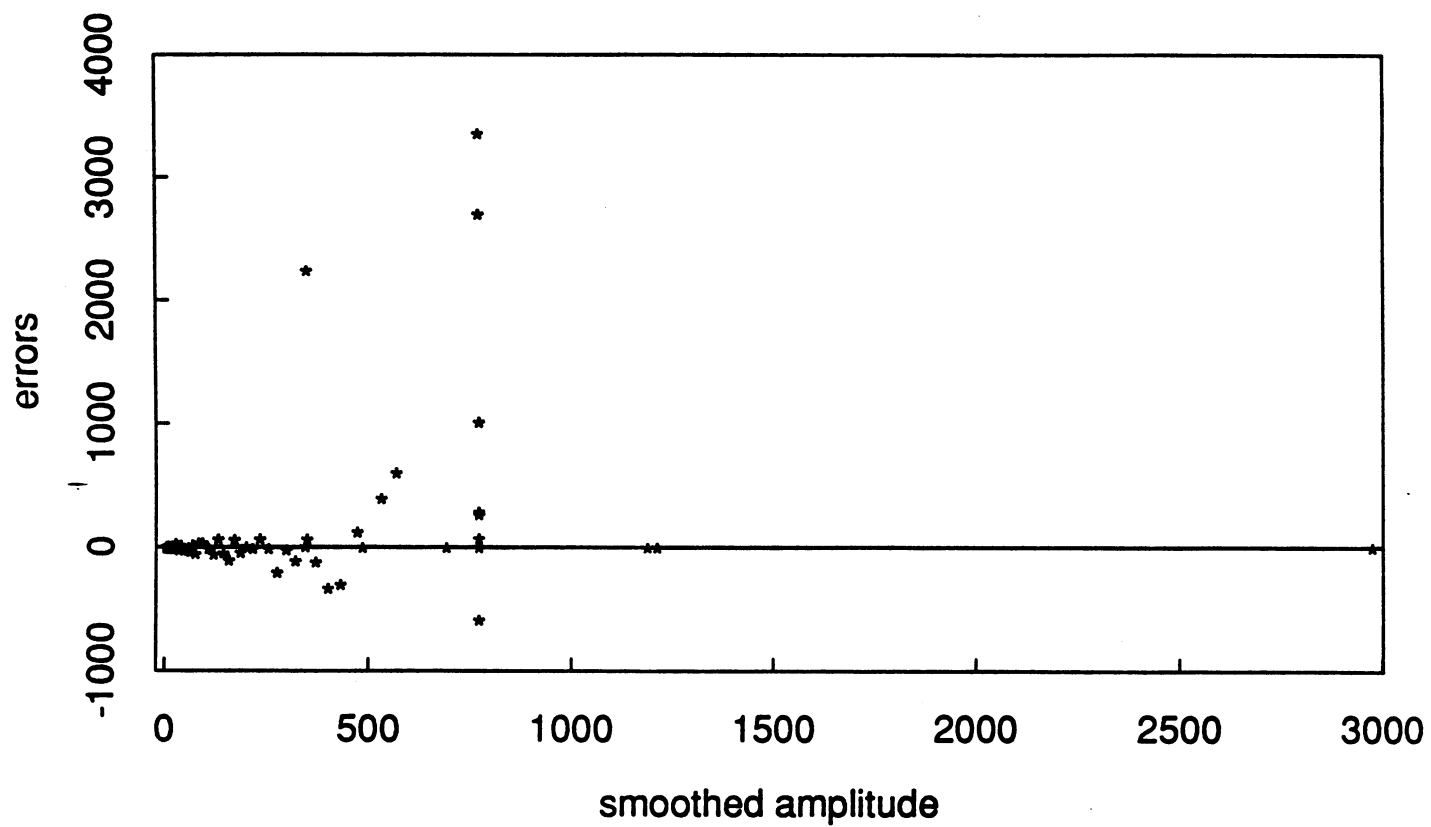
Taiwan Event - Amplitude Spectrum

frequency (cycles/second)
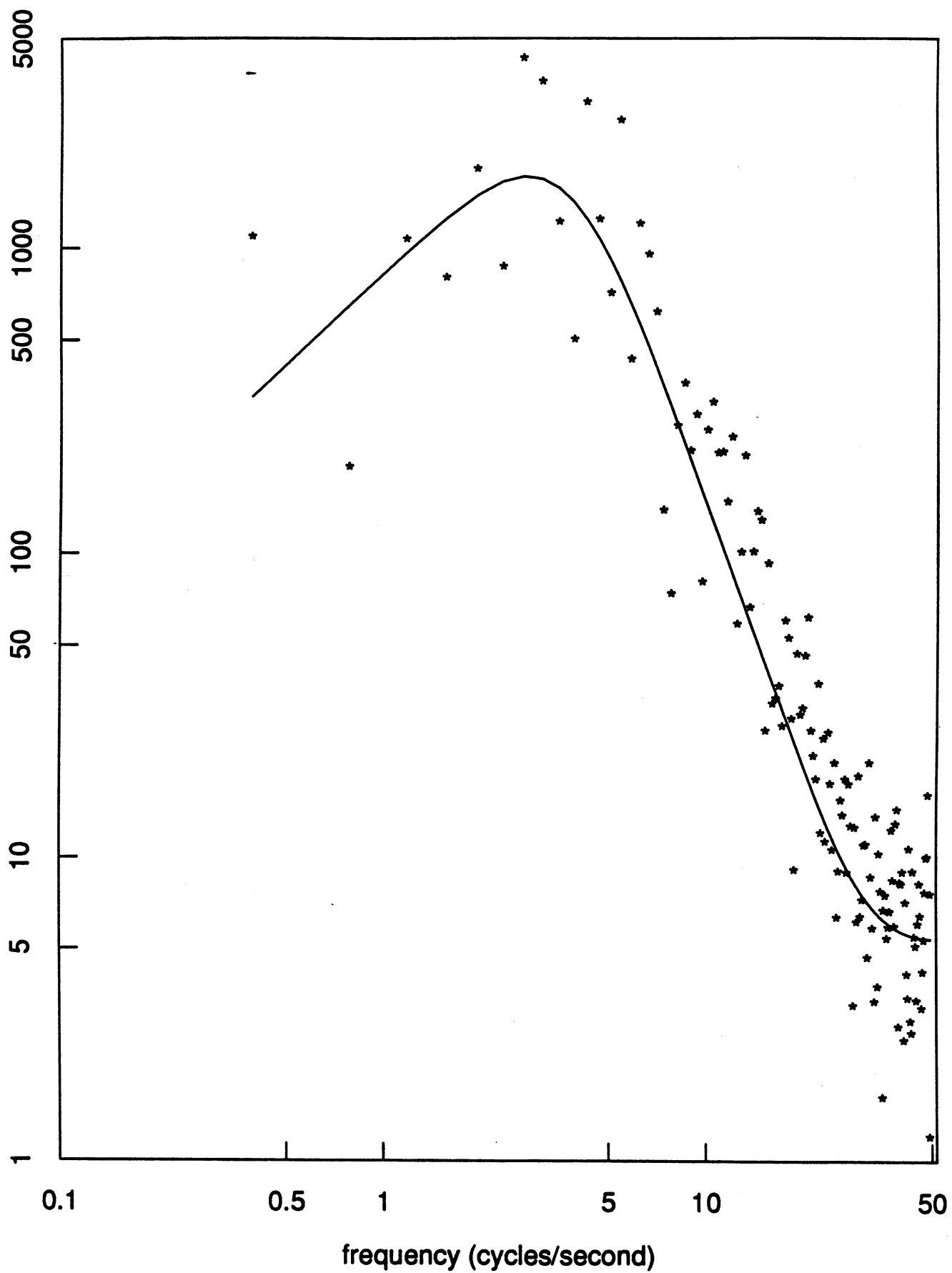
Taiwan Array and Event of 29 January 1981

distances from array center (meters)
* denotes sensor location
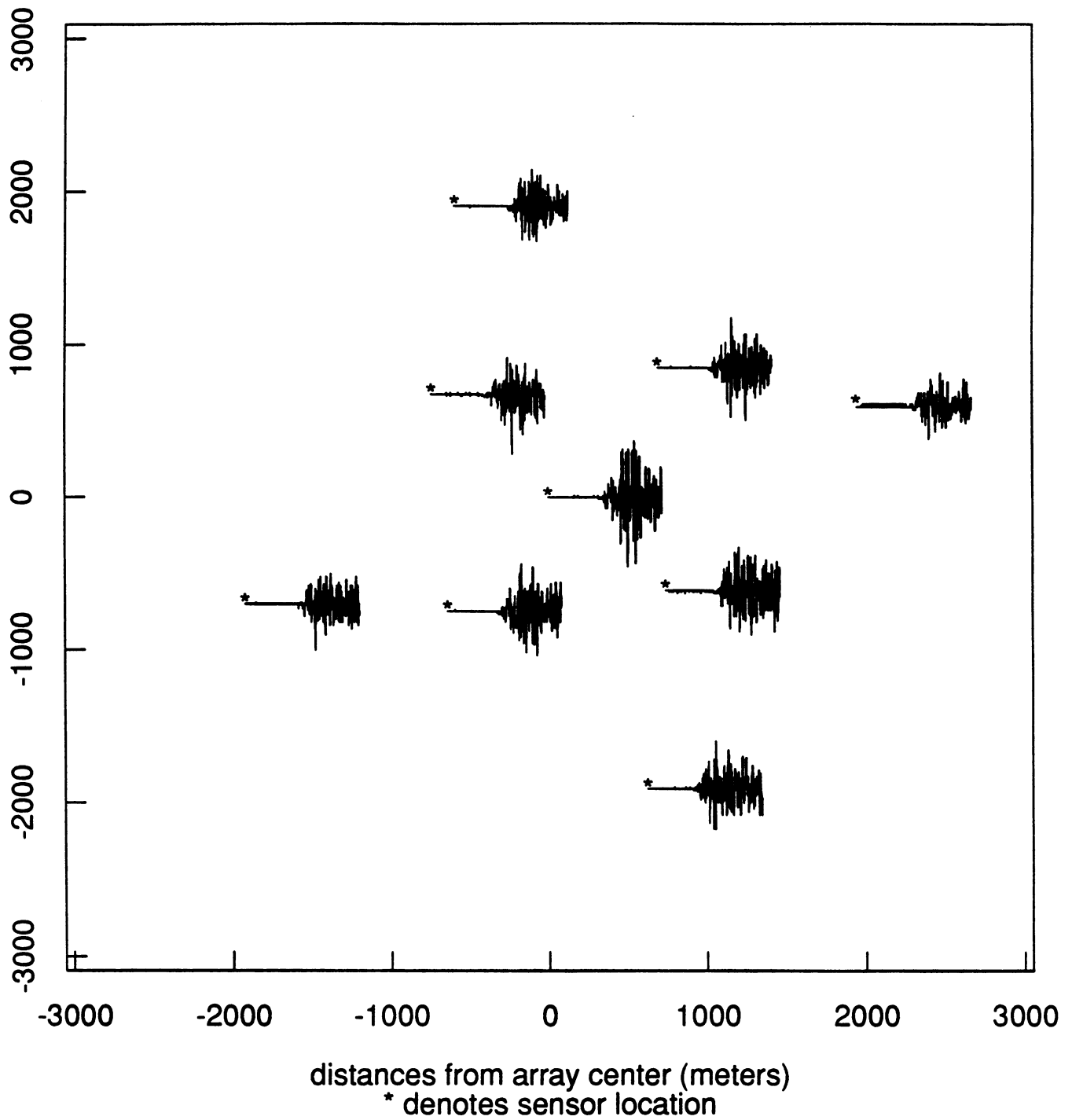
Frequency-Wavenumber Periodogram : Taiwan Event

wavenumber (cycles/km)

12
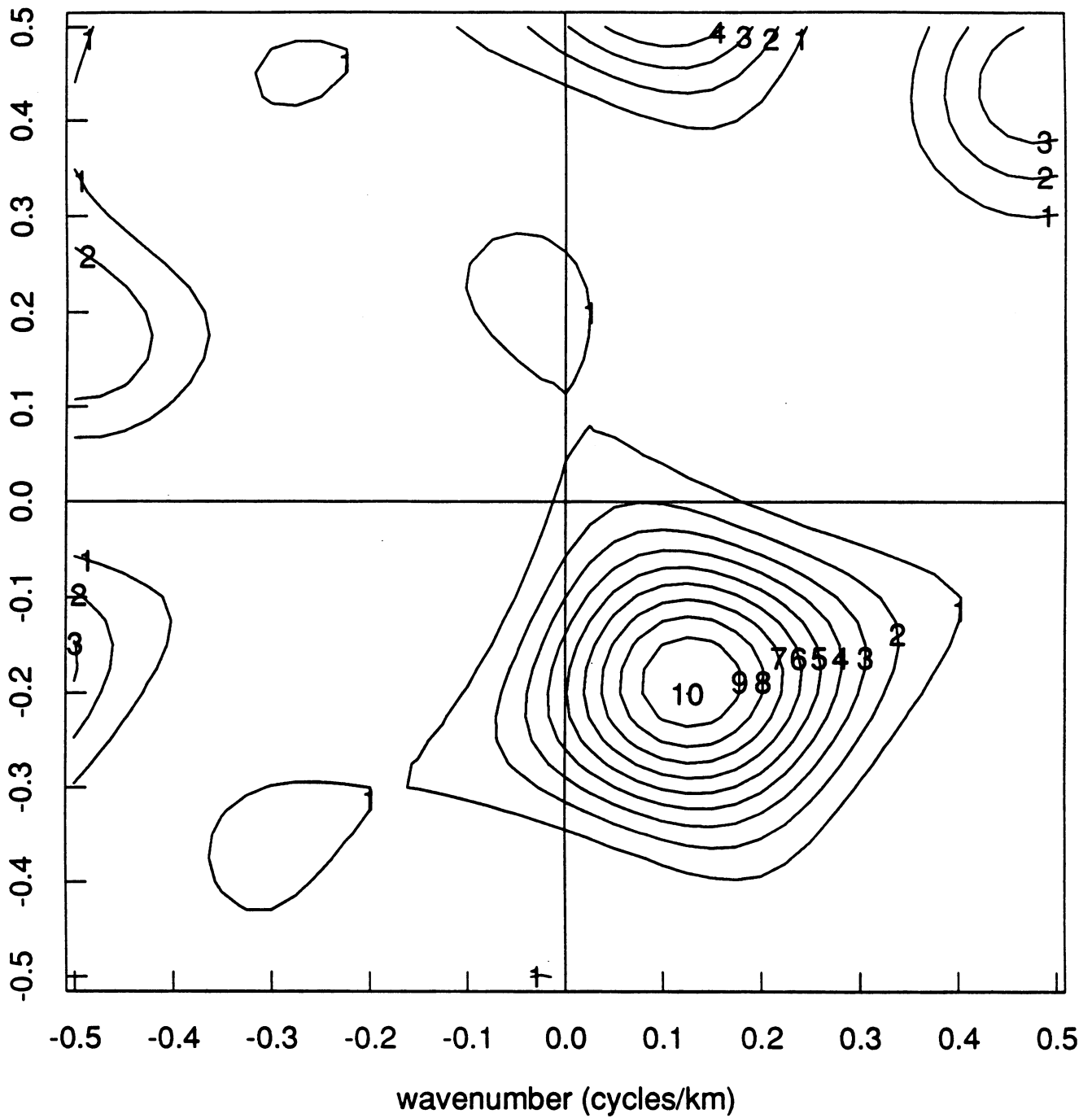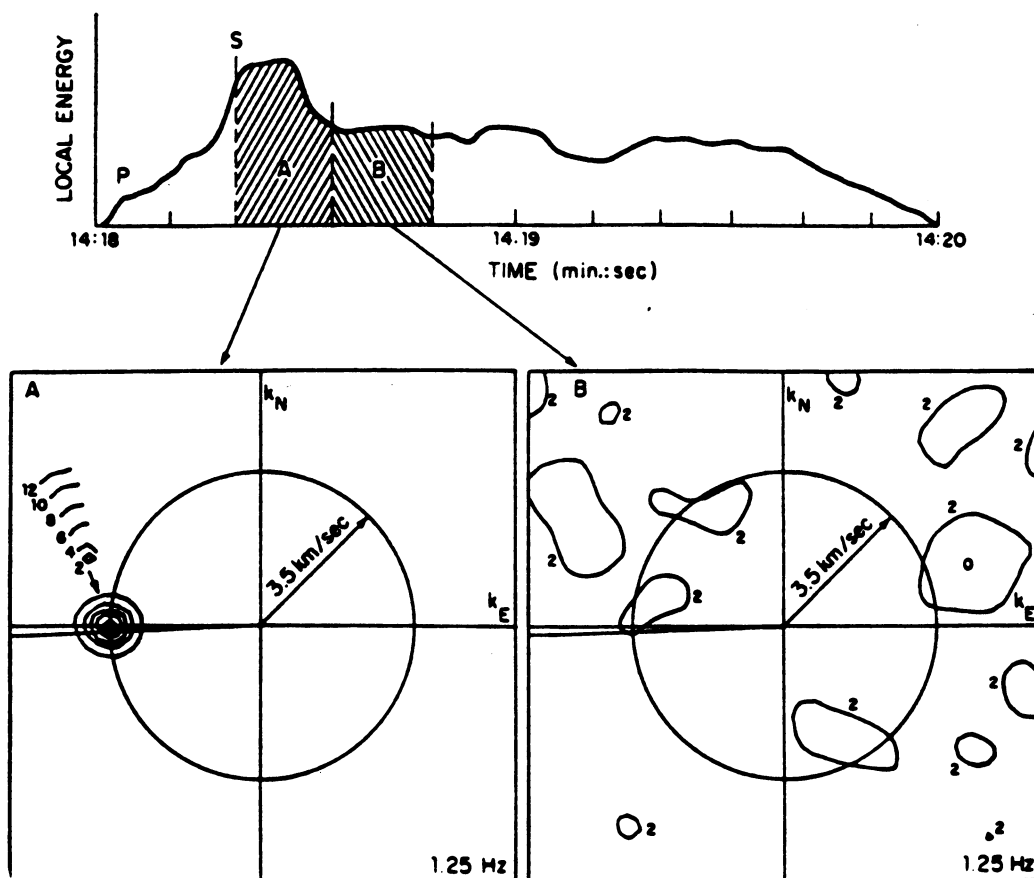
Autointensity Function - L10 as Pacemaker

lag (seconds)

Autointensity Function - L10 Bursting

lag (seconds)
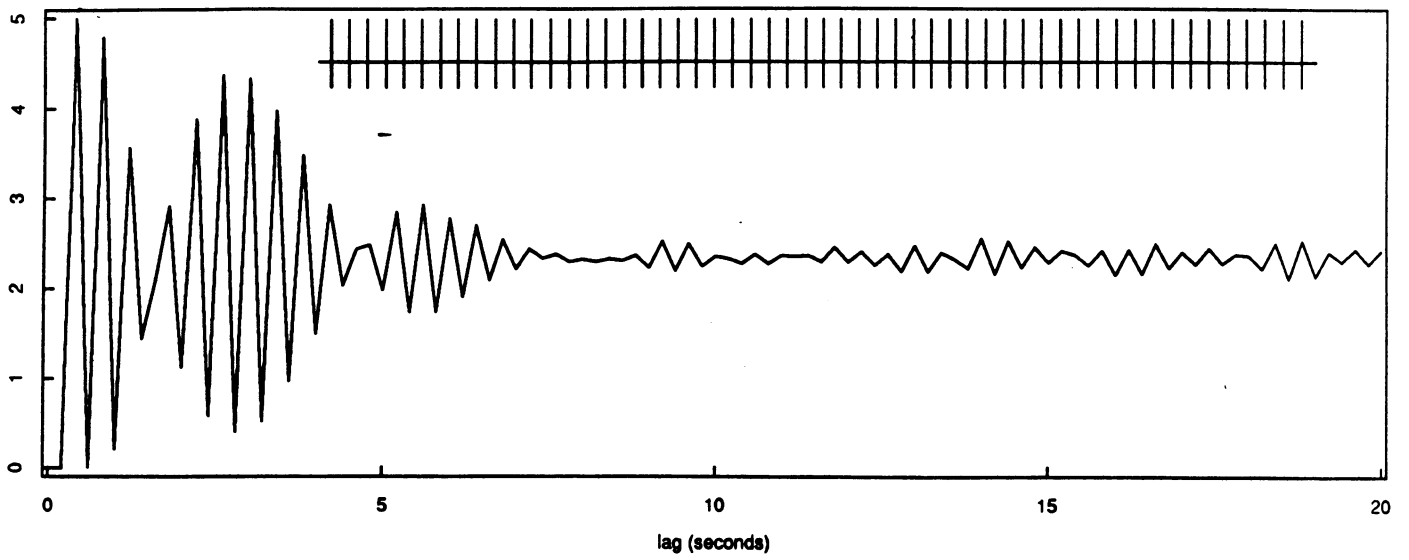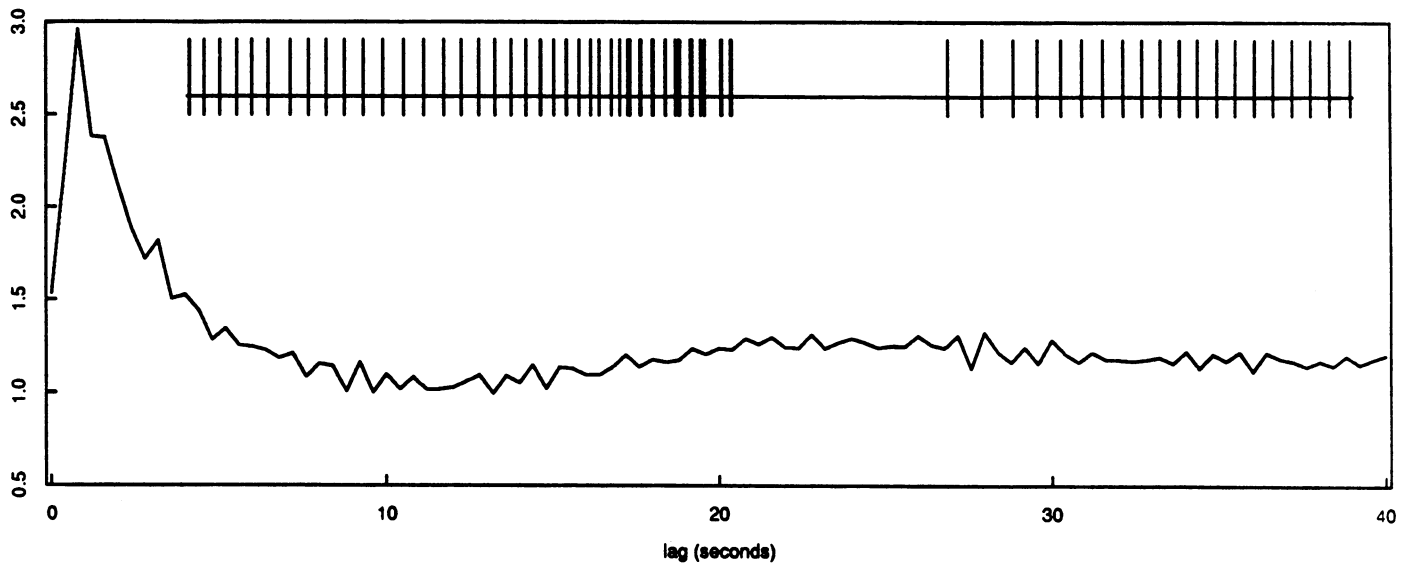
Autointensity Function - L10 Bursting with Accelerando

lag (seconds)

Crossintensity Function - L10 with L3

lag (seconds)
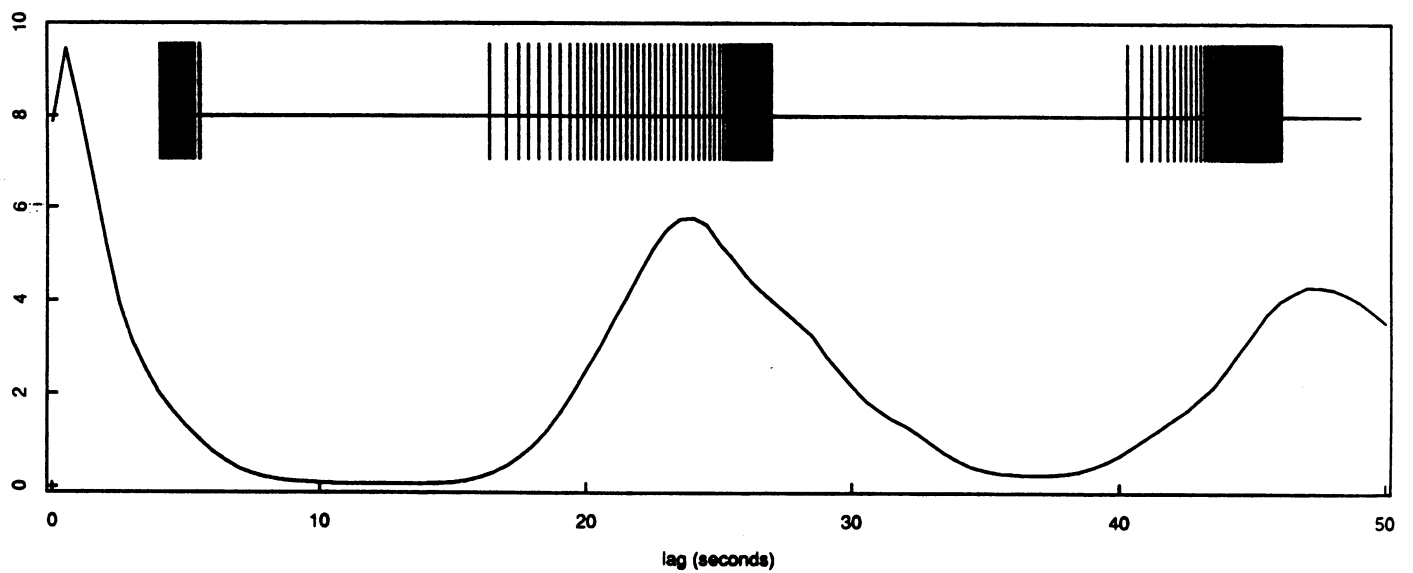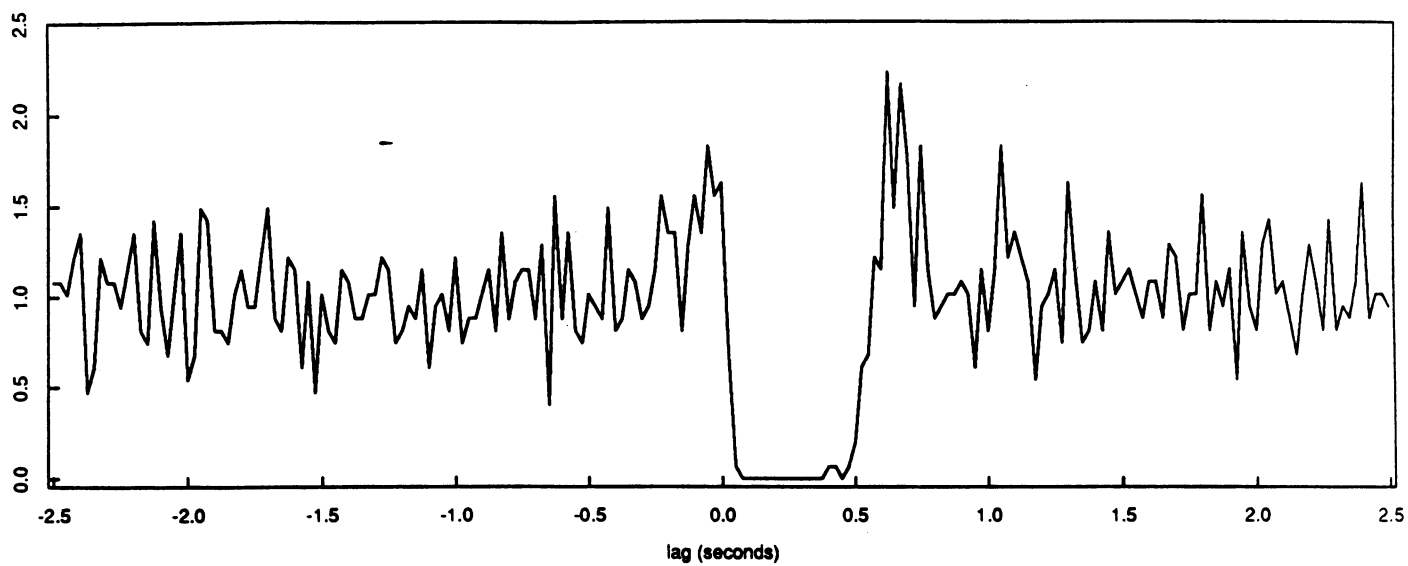
Crossintensity Function - RVP with R15
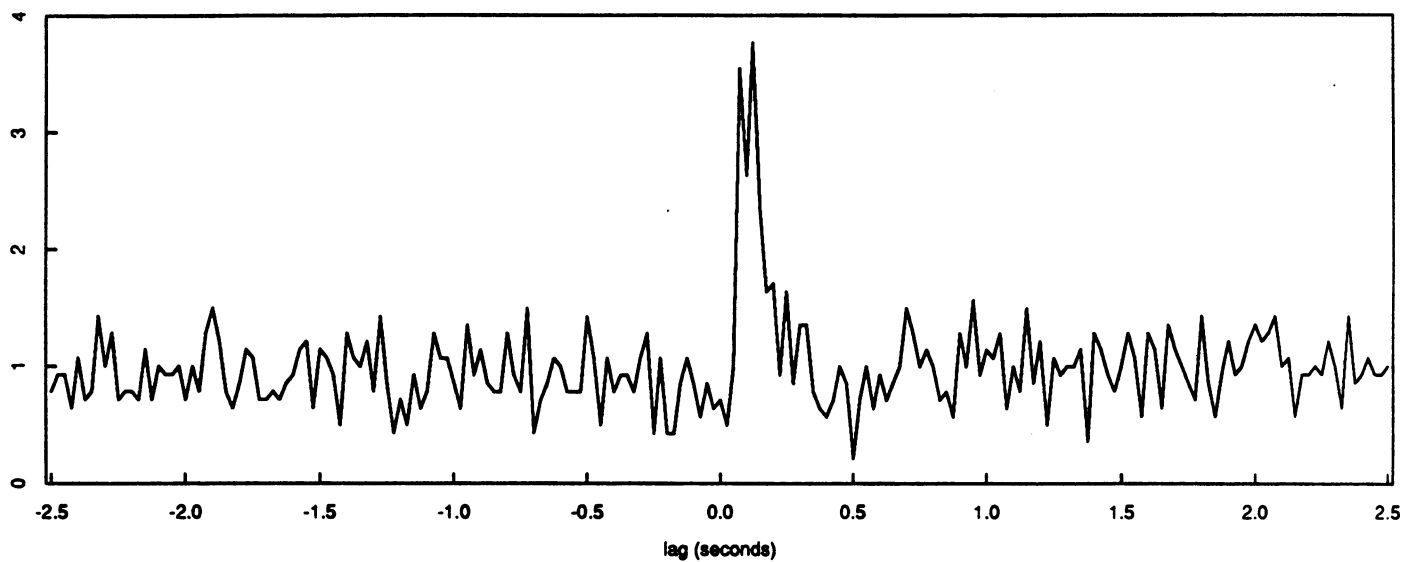
lag (seconds)

Crossintensity Function - L10 with L3

lag (seconds)

Coherence - L10 with L3

frequency (cycles/second)

Impulse Response

lag (seconds)
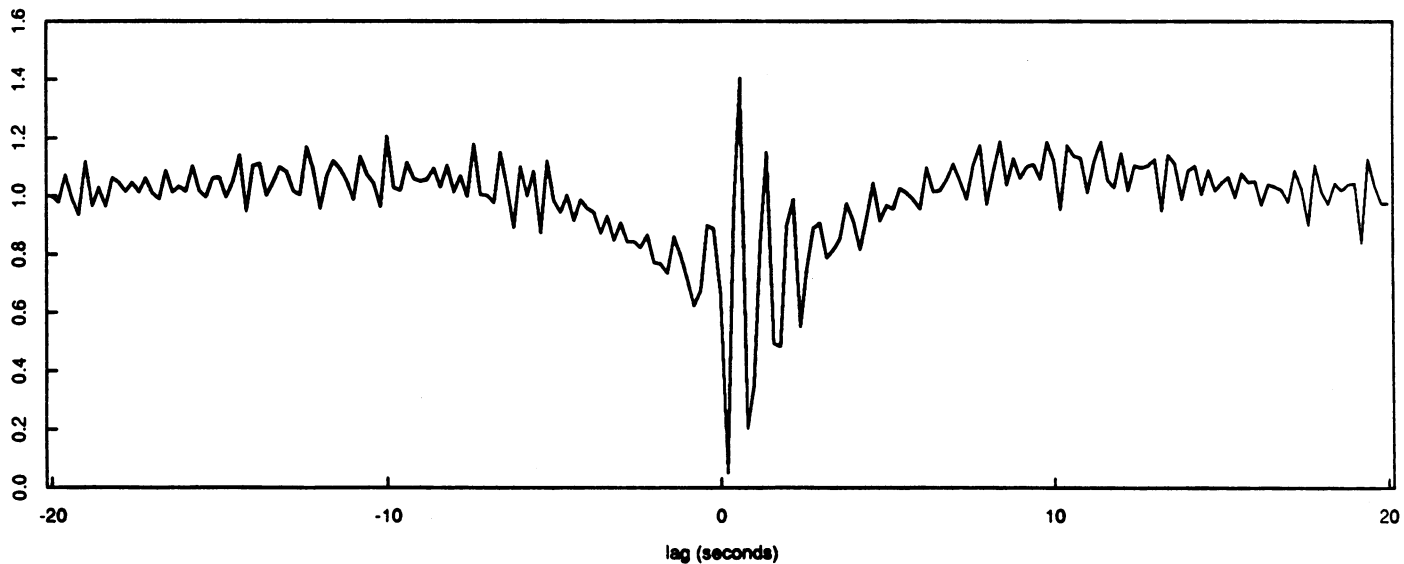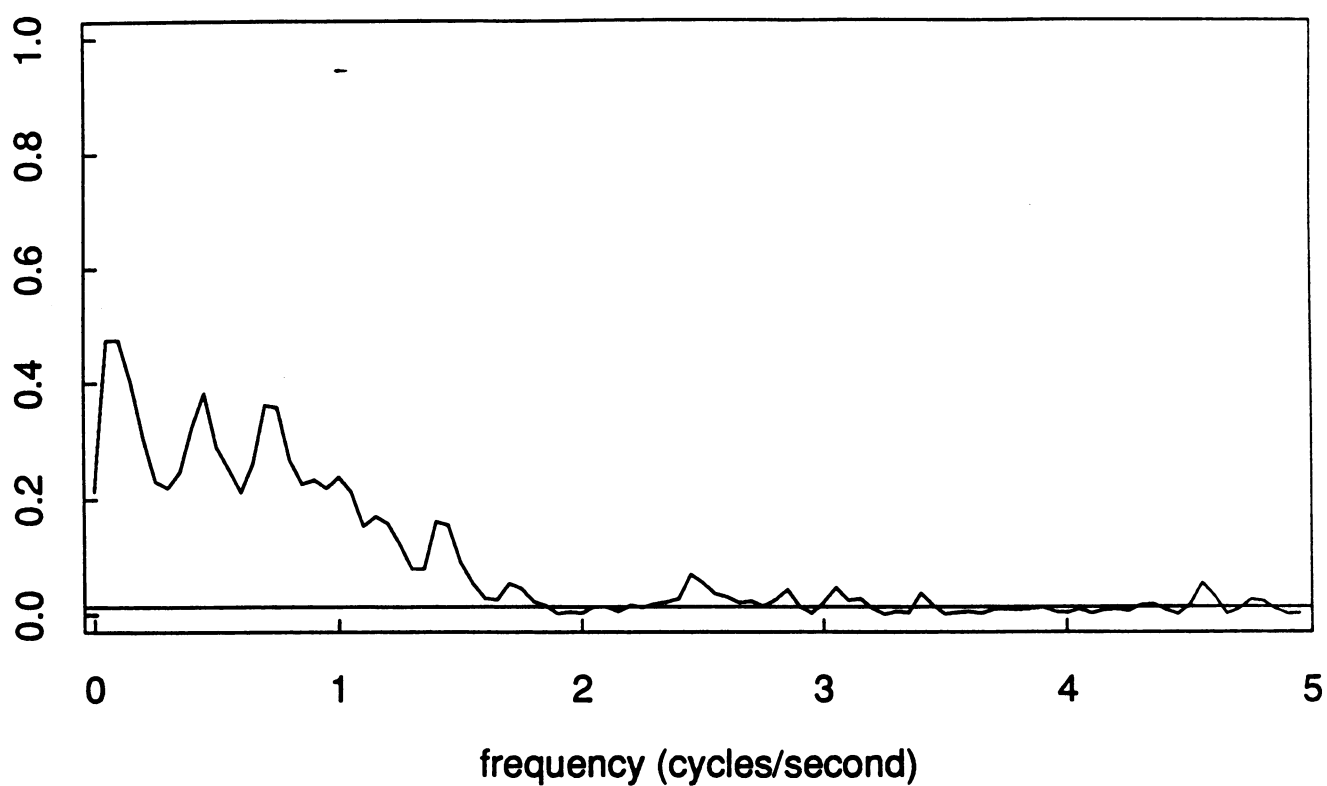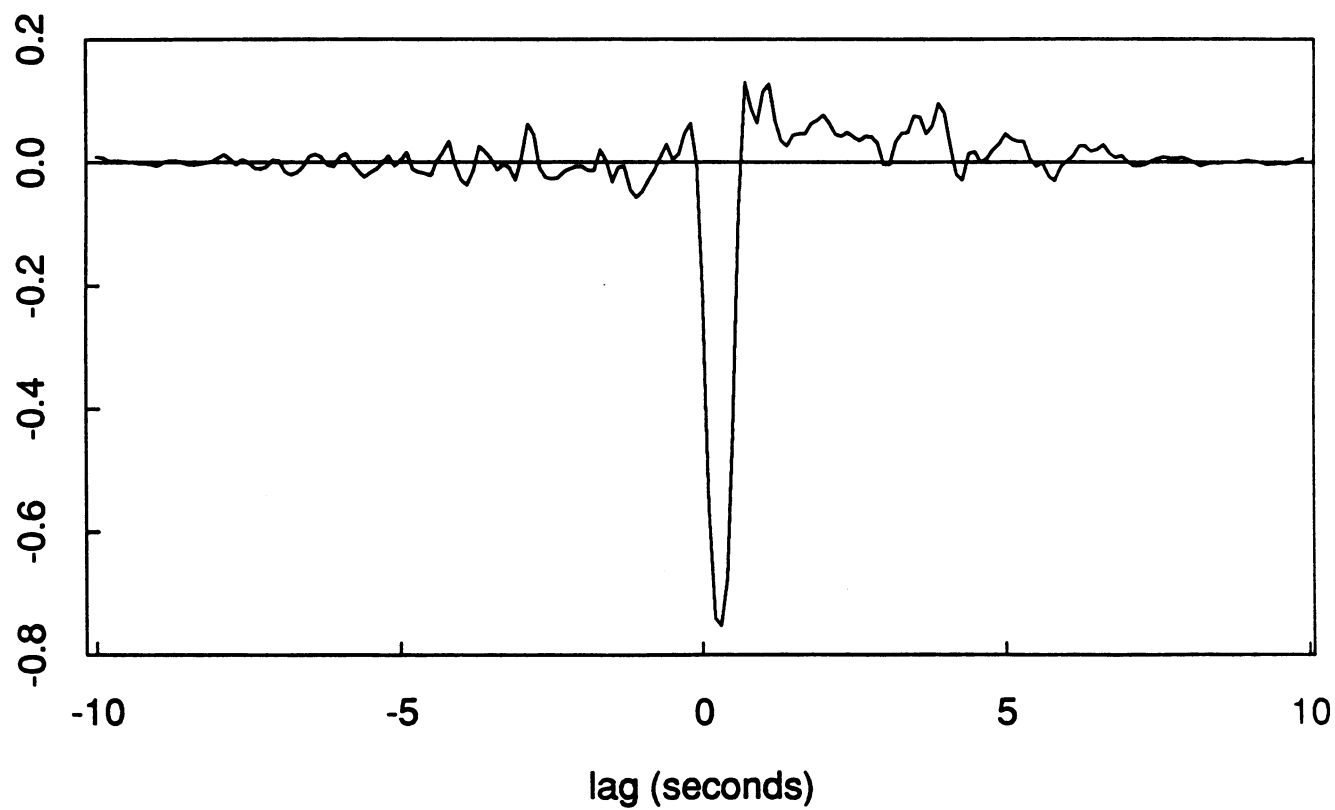
Coherence - RVP with R15

frequency (cycles/second)

Impulse Response

lag (seconds)
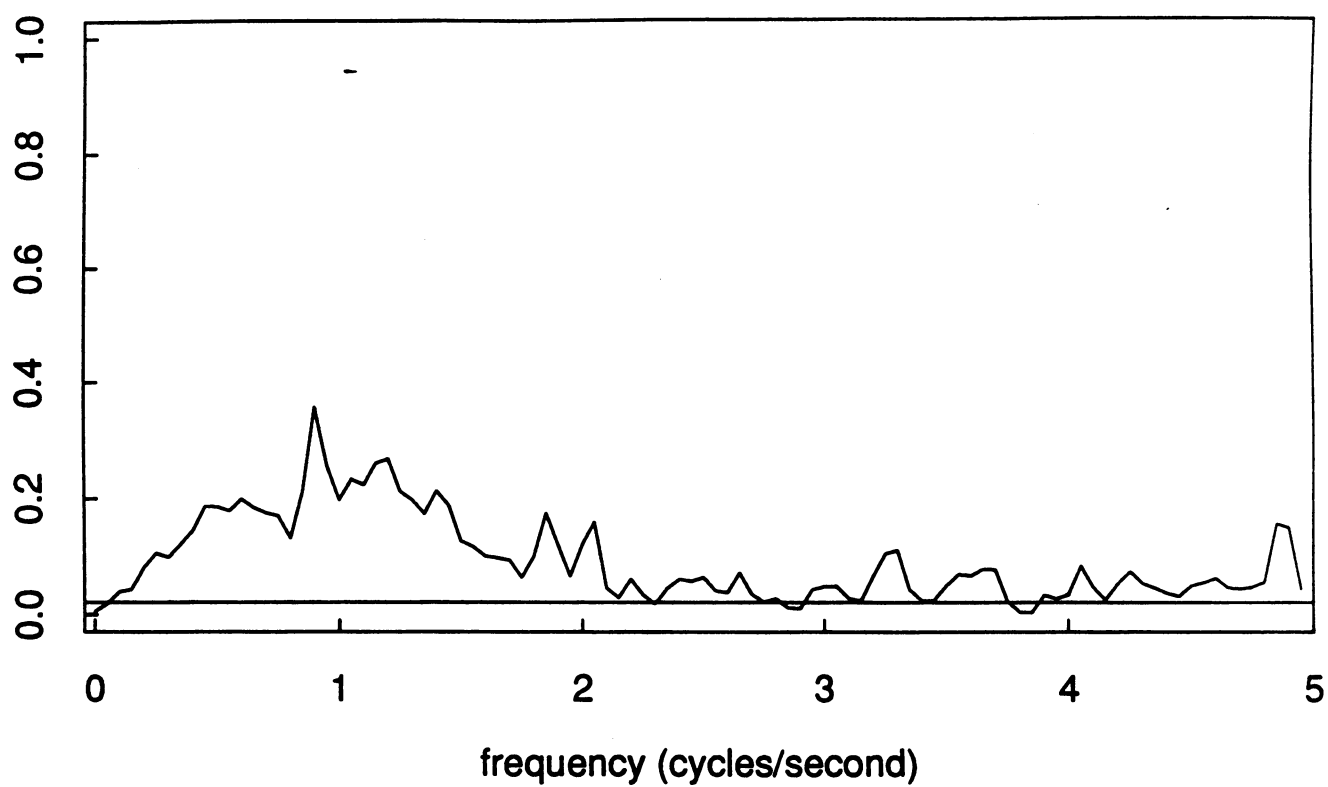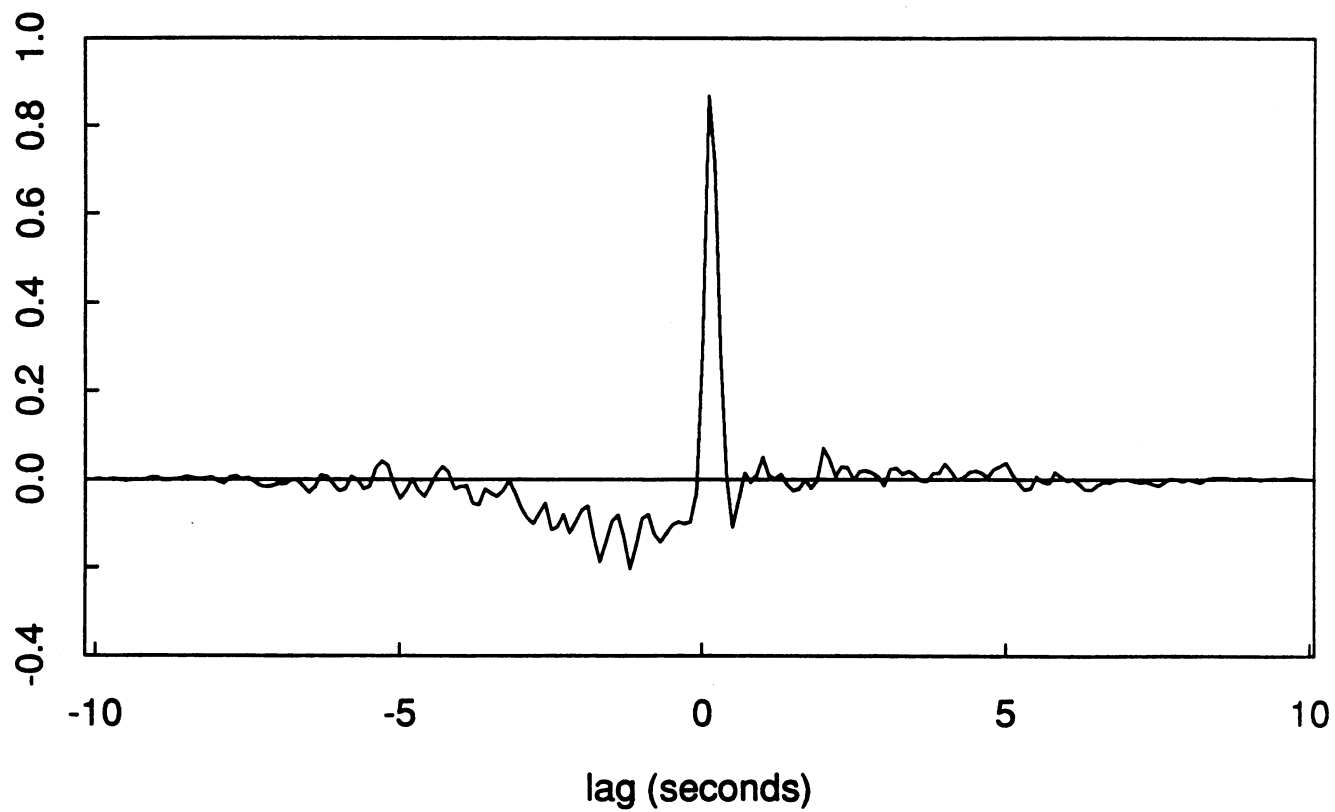
Coherence L10 and L2
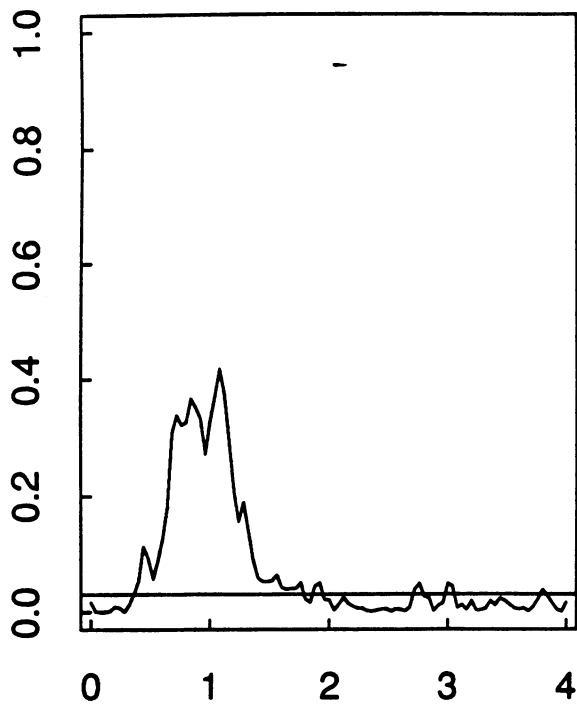
Coherence L2 and L3
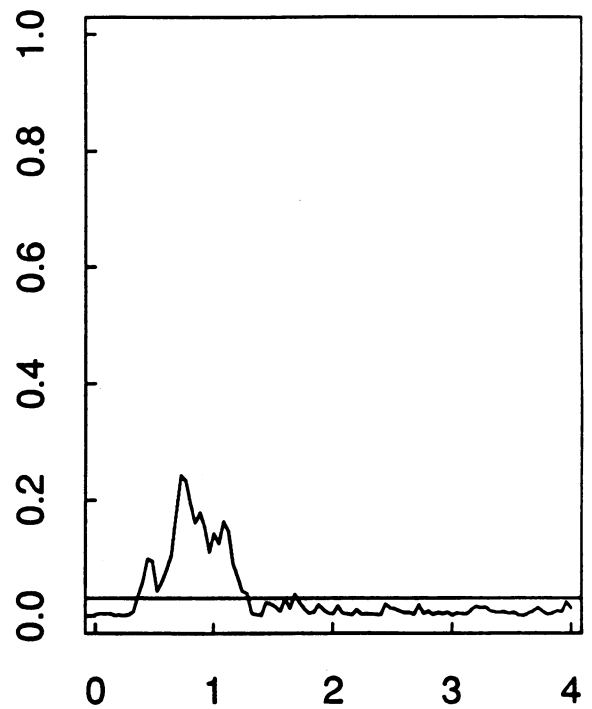
Coherence L10 and L3
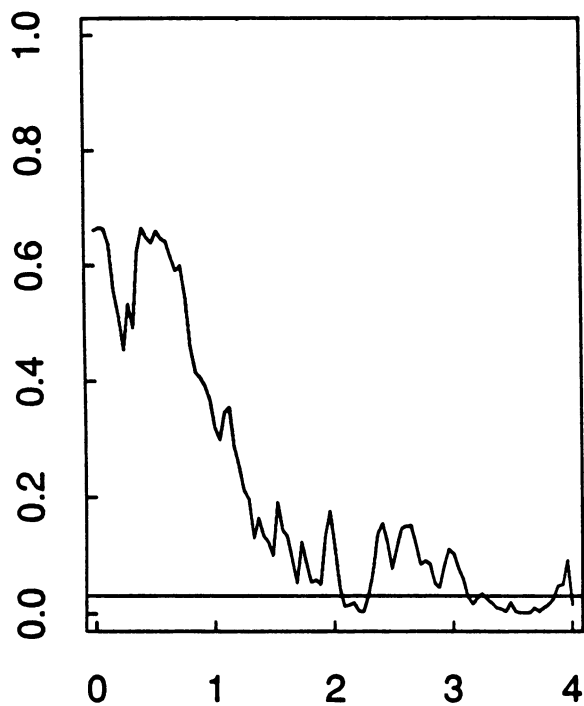
Partial coherence L2 and L3

frequency (Hz)
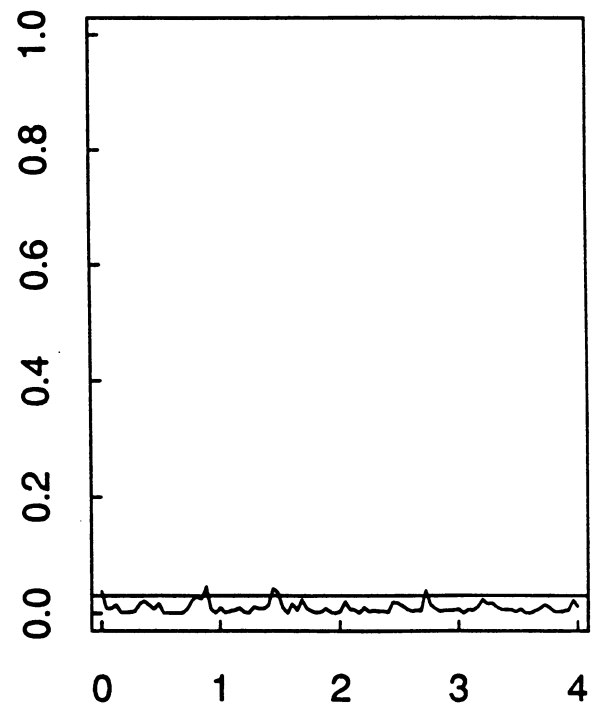
horizontal line : 95% null level

Neuron R2 - Noise Driven

time (seconds)

Summation Function

lag (seconds)
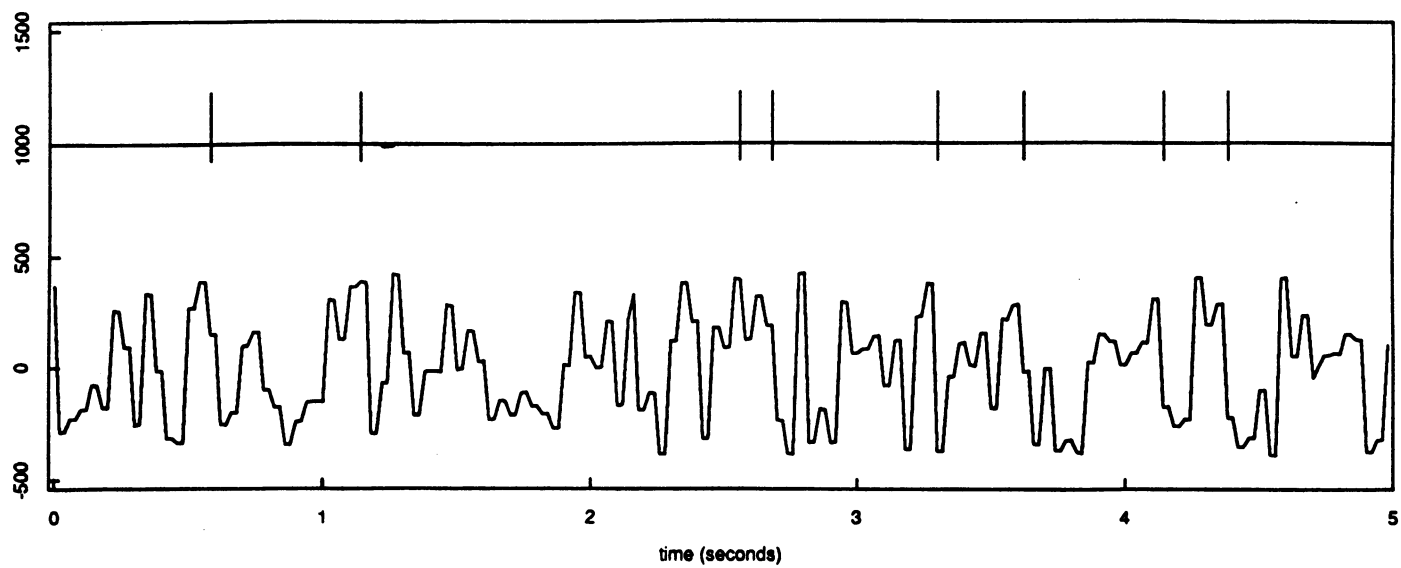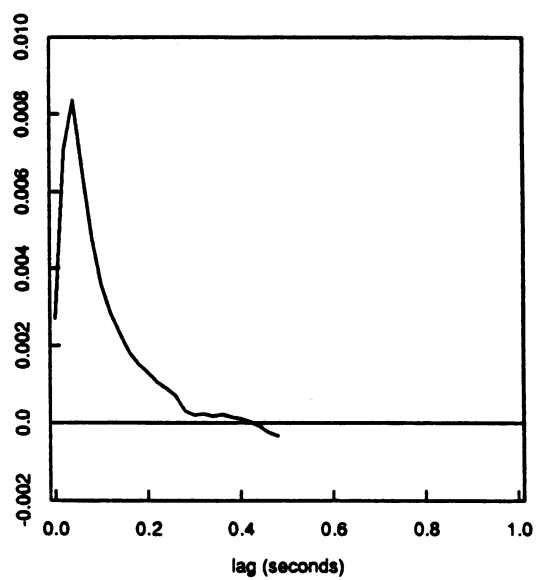
Empirical Probability of Firing

probability

linear predictor
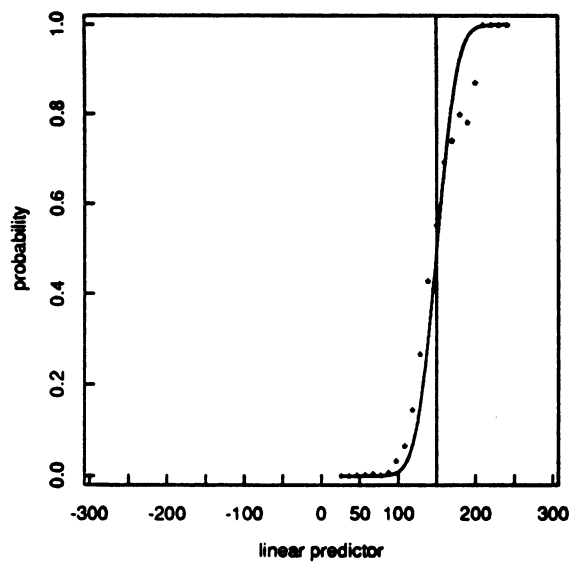
Rabbit Olfactory System - Responses at 4 by 4 Array
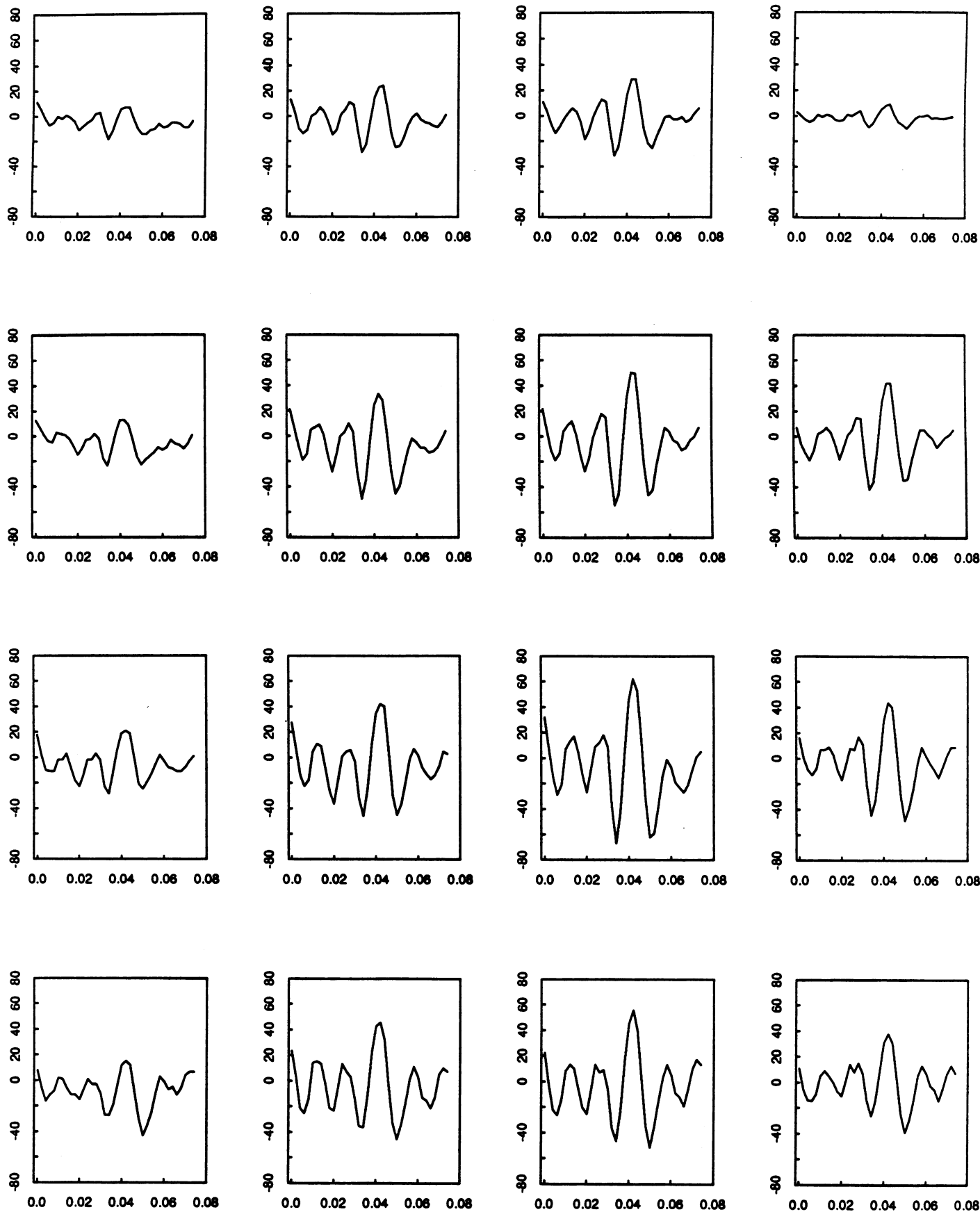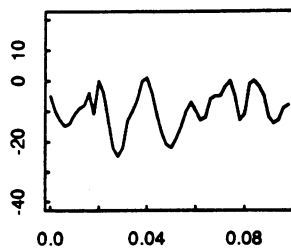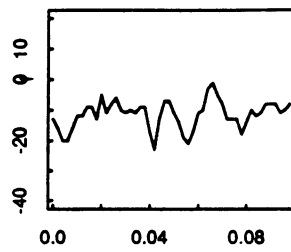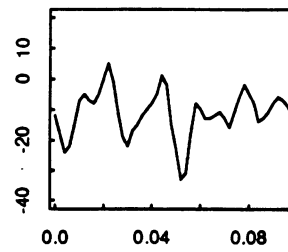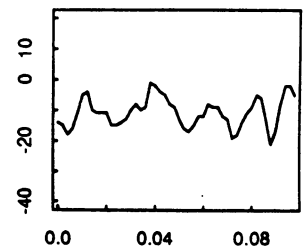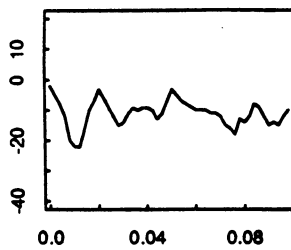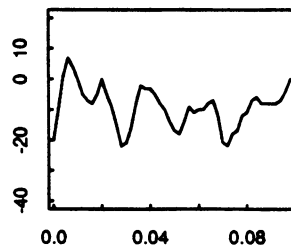
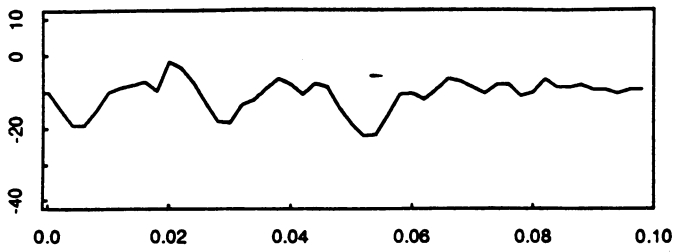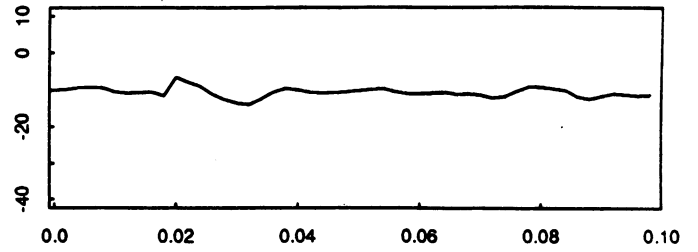# Individual Responses - Rabbit Pre-piriform Cortex
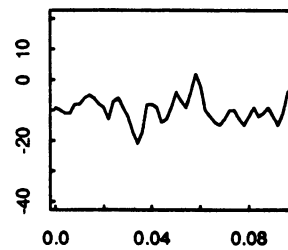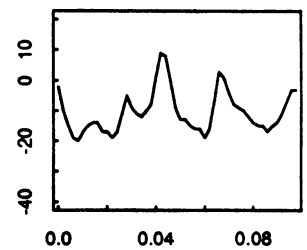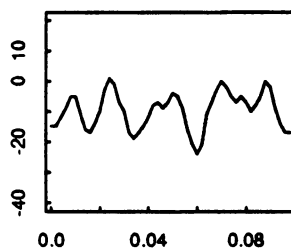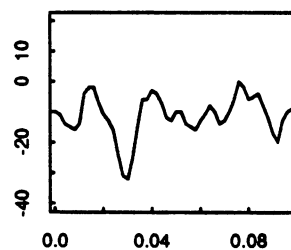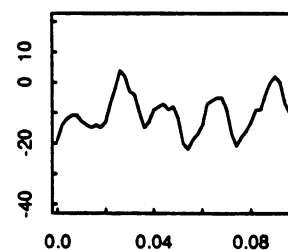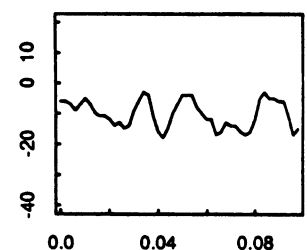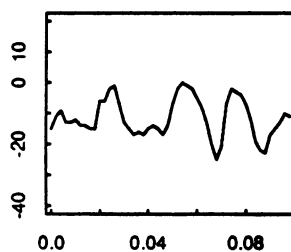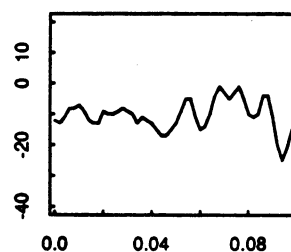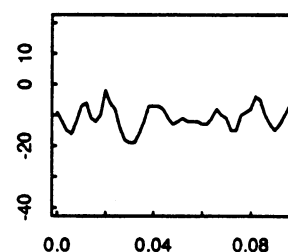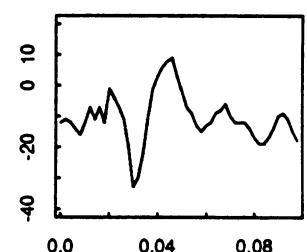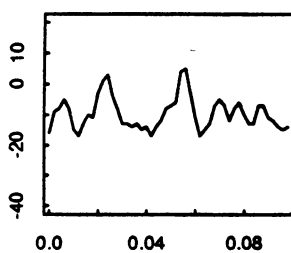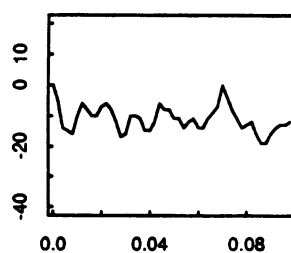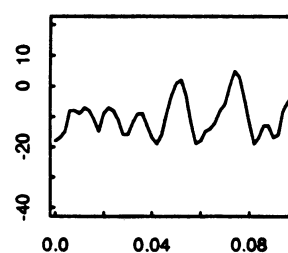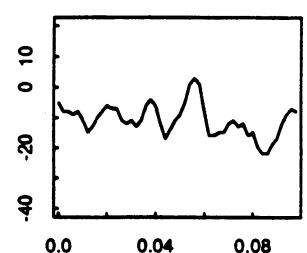
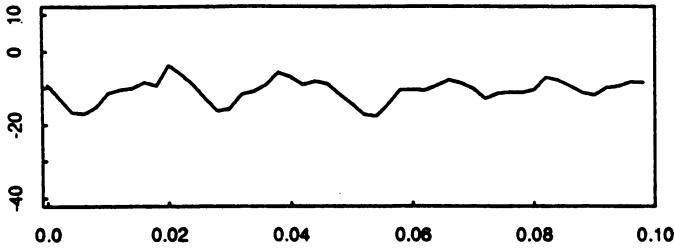# Average Evoked Responses - Several M's, Several Intensities

100% threshold - average of 3

89% threshold - average of 38

100% threshold - average of 5

111% threshold - average of 38

100% threshold - average of 10

122% threshold - average of 38

100% threshold - average of 20

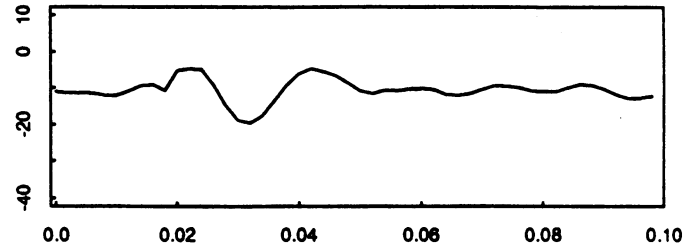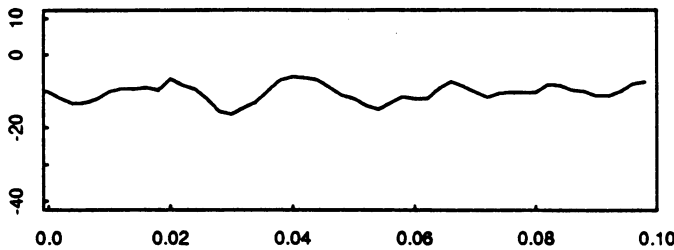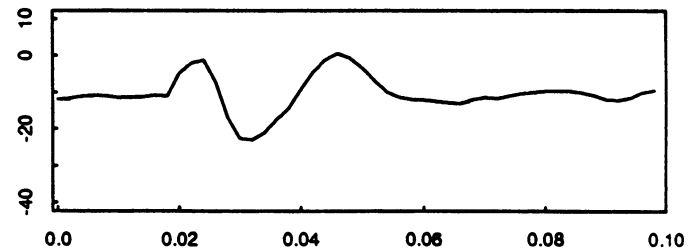133% threshold - average of 38

100% threshold - average of 38

144% threshold - average of 38
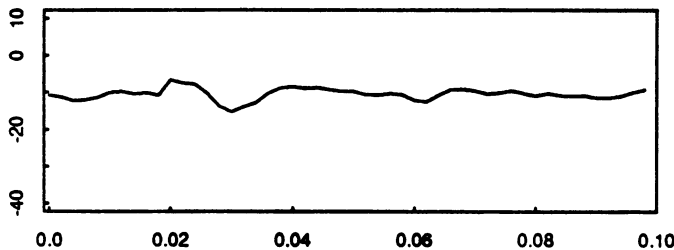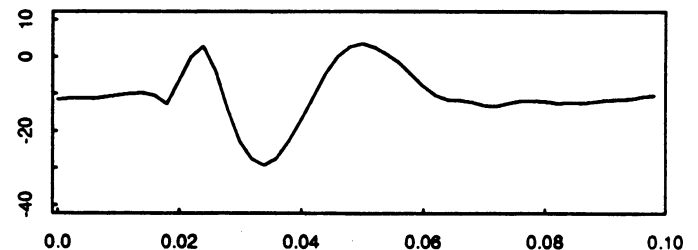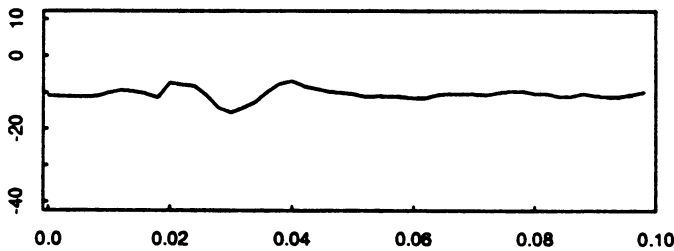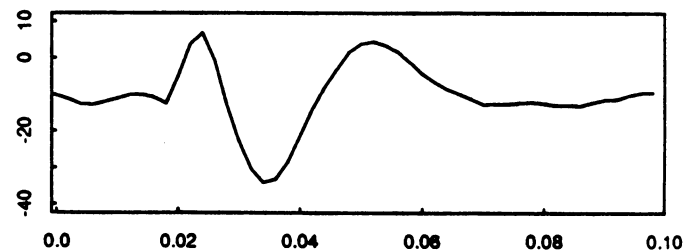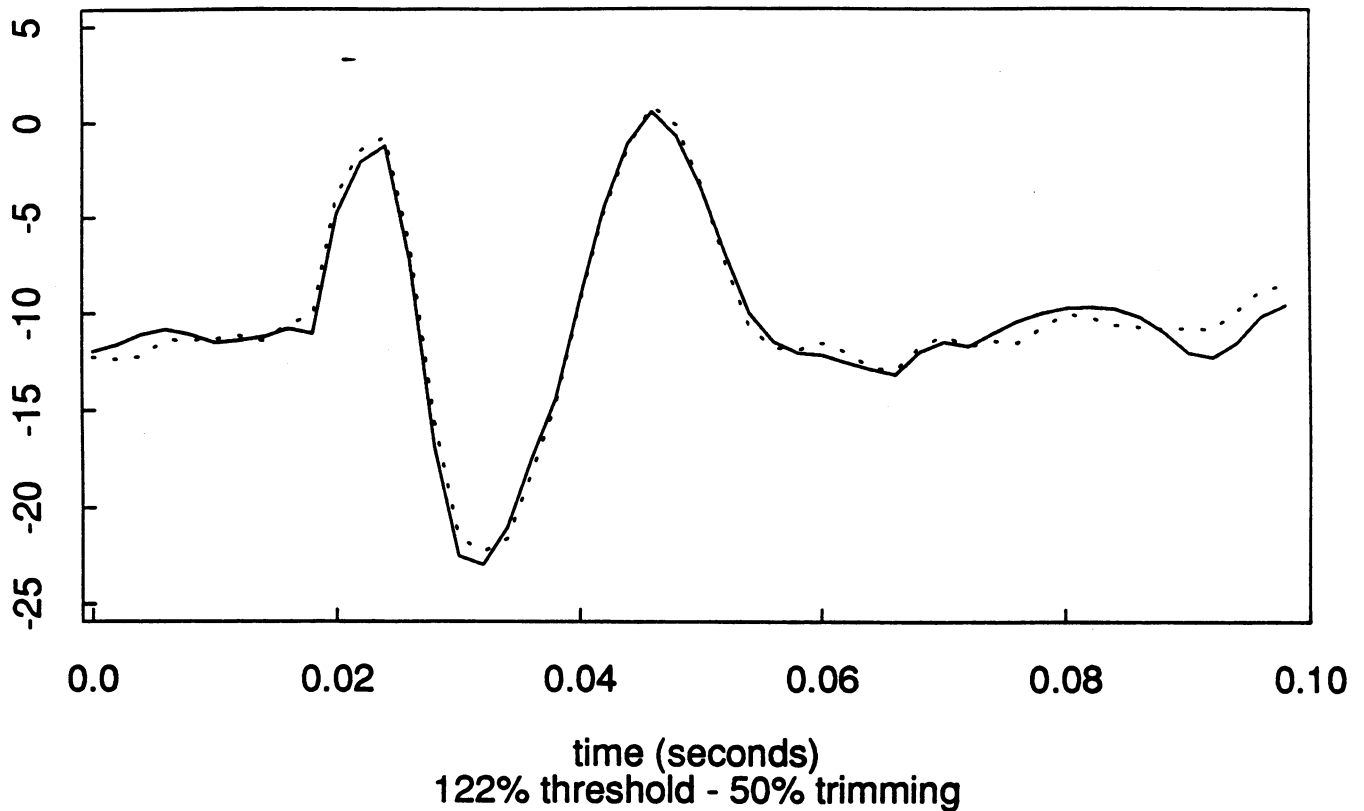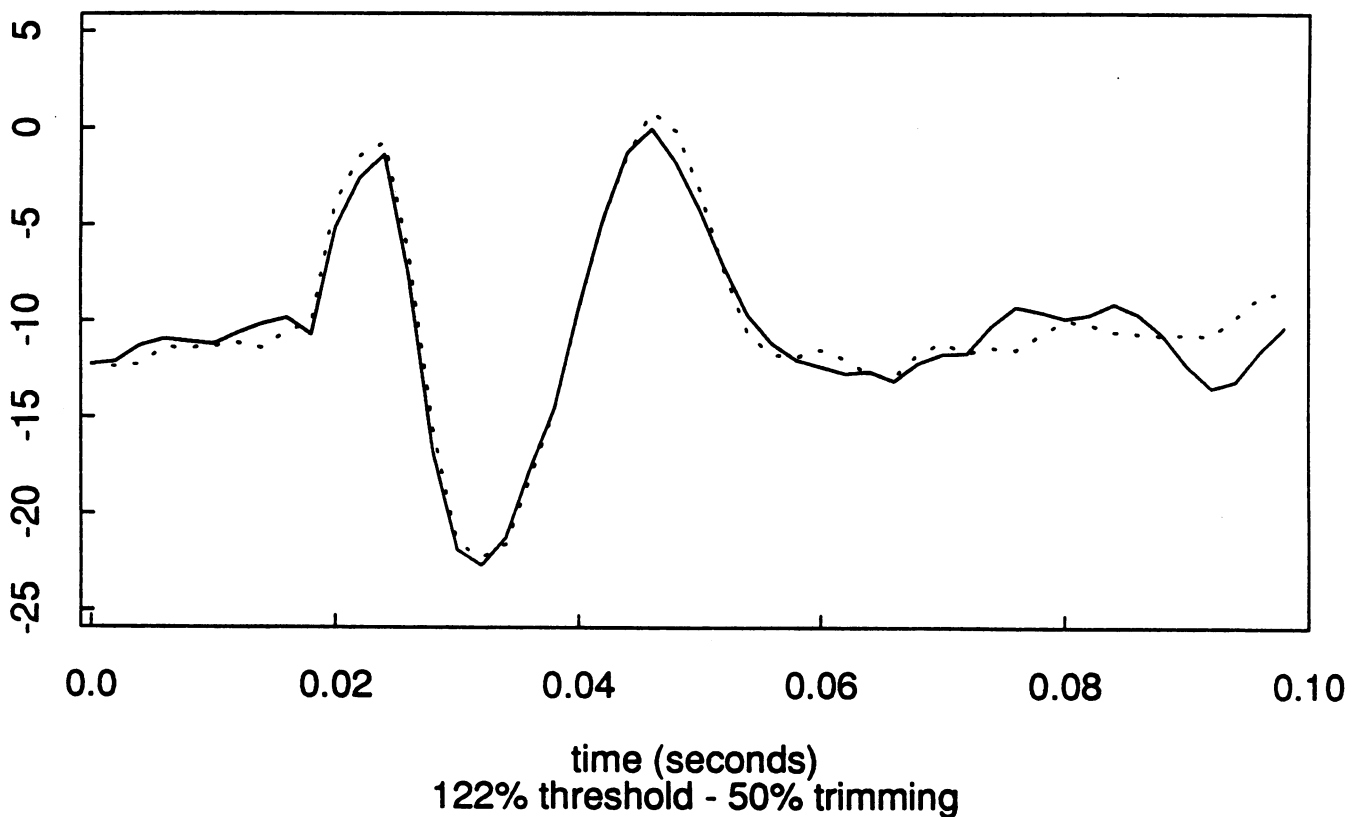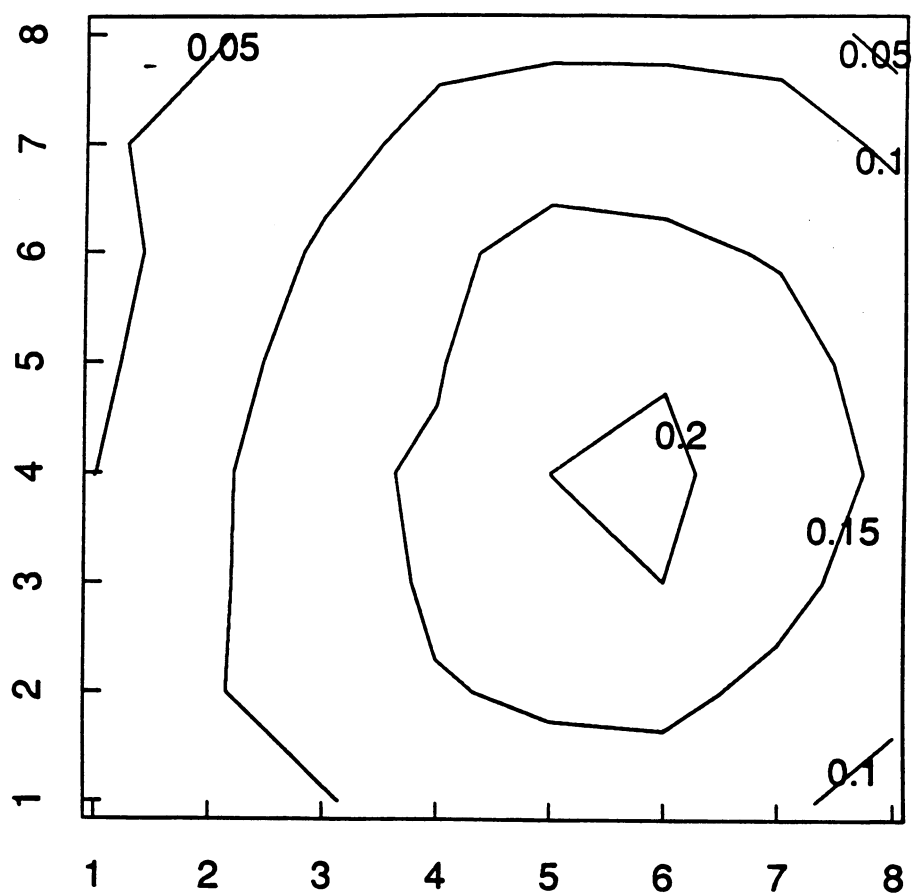
Average Evoked Response and Robust/Resistant Variant

time (seconds)
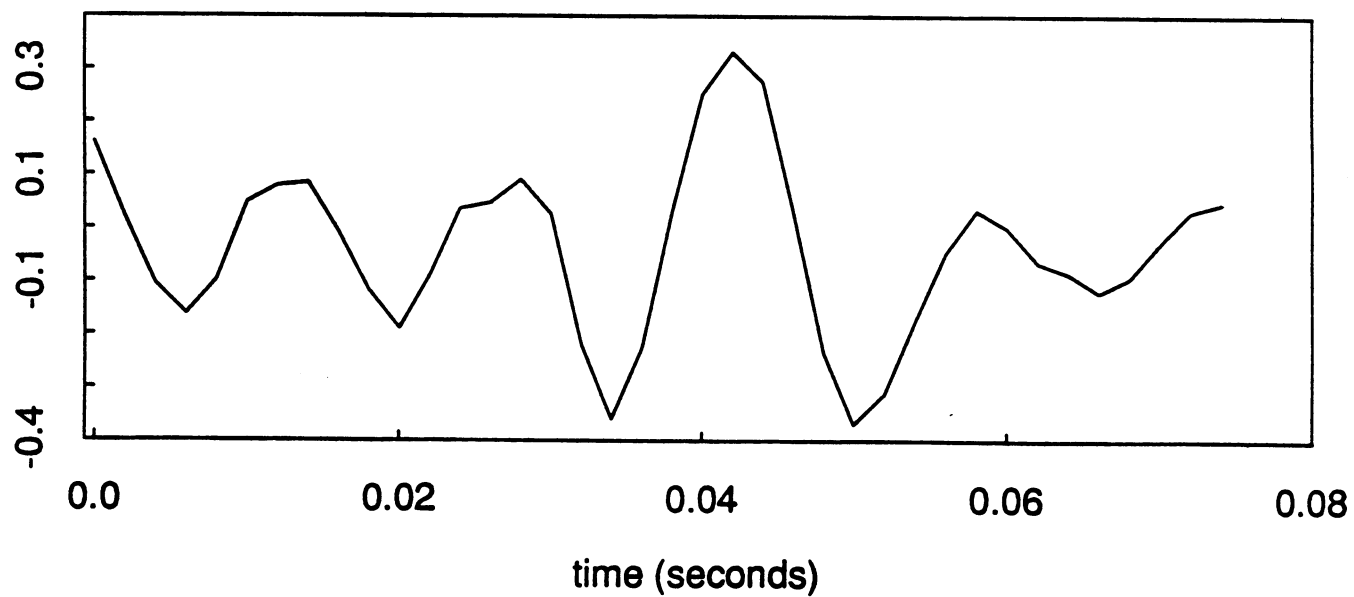122% threshold - 50% trimming

Running Trimmed Mean and Robust/Resistant Variant

time (seconds)
122% threshold - 50% trimming

Latent Spatial Component

Latent Temporal Component

time (seconds)

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Statistics Department
University of California, Berkeley

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